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GROUNDWATER MODELING REPORT

ASH POND NO. 1 COFFEEN POWER PLANT COFFEEN, ILLINOIS OUNDWATER MODELING REPORT

GROUNDWATER MODELING REPORT COFFEEN POWER PLANT ASH POND NO. 1

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ACRONYMS AND ABBREVIATIONS

EXECUTIVE SUMMARY

Ramboll Americas Engineering Solutions, Inc. (Ramboll) has prepared this Groundwater Modeling Report (GMR) on behalf of the Coffeen Power Plant (CPP), operated by Illinois Power Generating Company - IPGC (IPGC), in accordance with requirements of Title 35 of the Illinois Administrative Code (35 I.A.C.) Section (§) 845: Standards for the Disposal of Coal Combustion Residuals in Surface Impoundments (Part 845) (Illinois Environmental Protection Agency [IEPA], 2021). This document presents the results of predictive groundwater modeling simulations for proposed closure scenarios for the coal combustion residuals (CCR) management unit Ash Pond Number (No.) 1 (AP1 [(Vistra Identification [ID] No. 101, IEPA ID No. W1350150004-01, and National Inventory of Dams [NID] No. IL50722]). AP1 is a 23-acre, unlined surface impoundment (SI) used to manage CCR and non-CCR waste streams at the CPP. Its total storage capacity is approximately 300 acre-feet.

The CPP is located in Montgomery County, in central Illinois between the two lobes of Coffeen Lake (**Figure 1-1**), which was formed in 1963 by damming the McDavid Branch of the East Fork of Shoal Creek. Coffeen Lake encompasses approximately 1,100 acres and was created to provide a source of cooling water for the CPP. Coffeen Lake borders the CPP to the west, east, and south, and agricultural land is located to the north. Historically coal mines were operated at depth below the site. Mine shafts, processing facilities, and historic coal storage were located on the southern extent of the CPP, south of AP1. The CPP operated as a coal-fired power plant from 1964 until November 2019 and has five CCR management units, with AP1 being the subject of this GMR. Unlithified material present above the bedrock in the vicinity of the CPP was categorized into hydrostratigraphic units as part of the 2021 Hydrogeologic Site Characterization Reports (HCR; Ramboll, 2021a). In addition to the CCR, the hydrostratigraphic units occur in the following order (from ground surface downward) and include: Encourant presents the erast, of predictive groundweter modeling simulations for propossion execution frequent presents the results of predictive groundweter modeling simulations for propossion exerged and combustion resi

- **Upper Confining Unit (UCU):** Consists of the Loess Unit and the upper clayey portion of the Hagarstown Member which has generally lower vertical permeability. The UCU has been eroded east of AP1, near the Unnamed Tributary.
- **Uppermost Aquifer (UA):** The UA is the sandy portion of the Hagarstown Member which is classified as primarily sandy to gravelly silts and clays with thin beds of sands. Similar to the Loess Unit, the Hagarstown is absent in some locations near the Unnamed Tributary.
- **Lower Confining Unit (LCU):** Comprised of the Vandalia Member, Mulberry Grove Member, and Smithboro Member. These units include a sandy to silty till with thin, discontinuous sand lenses, a discontinuous and limited extent sandy silt which has infilled prior erosional features, and silty to clayey diamicton, respectively.
- **Deep Aquifer (DA):** Sand and sandy silt/clay units of the Yarmouth Soil, which include accretionary deposits of fine sediment and organic materials, typically less than five feet thick and discontinuous across the CPP.
- **Deep Confining Unit (DCU):** Comprised of the Banner Formation and generally clays, silts, and sands. The Lierle Clay Member is the upper layer of the Banner Formation which was encountered at the CPP.

Flow of groundwater from central portions of the CPP to Coffeen Lake or the Unnamed Tributary through the UA are the primary pathways for contaminant migration. Groundwater elevations are

primarily controlled by surface topography, geologic unit topography, and water levels within Coffeen Lake and the Unnamed Tributary. A groundwater divide trending north-south is observed running through the approximate center of the CPP. Phreatic surfaces or water elevations within the SI are generally consistent and have not been observed to fluctuate with groundwater elevations, indicating limited hydraulic connection with the SI.

The conceptual site model (CSM) for modeling the groundwater at the CPP is as follows:

- Most hydrostratigraphic layers are laterally continuous across the area. The flat to gently rolling uplands are dissected by deeply incised streams (into the materials of the UCU, UA, and LCU) that are tributaries to river systems in the area. Coffeen Lake was created by damming one of these tributary streams for use by the CPP.
- The UA is separated from the bottom of the AP1 by a minimum of 10 feet of low-permeability glacial till that comprises the UCU. Erosion caused by incised streams has occurred along the northeast corner of AP1 which likely results in ash being in contact with the UA.
- Surface recharge and groundwater migrate vertically through the low permeability sediments of the UCU. Groundwater migrates horizontally through the higher permeability sediments of the UA.
- Groundwater elevations and lake elevations indicates groundwater flows into Coffeen Lake from the UA.
- AP1 is constructed such that the earthen berm and base are in contact with the UCU with exception of limited areas in the northeast of the SI where the UCU and UA have been eroded and the berm and base are in contact with the LCU.
- The stage within AP1 is managed with minimal (less than 3 feet) variability throughout the year.

A review and summary of data collected from 2015 through 2021 for parameters with groundwater protection standards (GWPS) listed in 35 I.A.C. § 845.600 is provided in the HCR (Ramboll, 2021a). Concentration results presented in the HCR and summarized in the History of Potential Exceedances (Ramboll, 2021b) are considered potential exceedances because the methodology used to determine them is proposed in the Statistical Analysis Plan (Appendix A to the Groundwater Monitoring Plant [GMP], Ramboll 2021c), which has not been reviewed or approved by IEPA at the time of submittal of the Part 845 operating permit application. The following constituents with potential exceedances of the GWPS listed in 35 I.A.C. § 845.600 were identified: boron, cobalt, pH, sulfate, and total dissolved solids (TDS) (Ramboll, 2021b) at AP1. Most hydrostratigraphic layers are laterally continuous across the area. The flat to gen rolling uplands are aliseseded by deeply incided streams (into the materials of the UCU, that are tristochers to the very systems in

A Technical Memorandum (**Attachment A**) was prepared by Geosyntec Consultants, Inc. (Geosyntec, 2022a), *Draft Evaluation of Potential Groundwater Protection Standard Exceedances, Coffeen Ash Pond No.1, Coffeen Illinois*, to further evaluate potential GWPS exceedances. The results of the evaluation demonstrated that the potential GWPS exceedances of cobalt in well G314 and pH in well G312 are not related to AP1 based on several lines of evidence presented in the Technical Memorandum.

Statistically significant correlations between sulfate concentrations and concentrations of TDS identified as potential exceedances of the GWPS indicate sulfate is an acceptable surrogate for TDS in the groundwater model. Concentrations of TDS are expected to change along with model predicted sulfate concentrations. A potential exceedance of boron was observed at one

monitoring well, G313, which also has potential exceedances of both sulfate and TDS. Similar source and behavior in the groundwater system would be expected among boron, sulfate, and TDS at UA monitoring well G313, and boron concentrations are expected to change along with model predicted sulfate concentrations.

It was assumed that sulfate would not significantly sorb or chemically react with aquifer solids (distribution coefficient [Kd] was set to 0 milliliters per gram [mL/g]) which is a conservative estimate for predicting contaminant transport times in the model. Boron, sulfate, and TDS transport is likely to be affected by both chemical and physical attenuation mechanisms (*i.e.,* adsorption and/or precipitation reactions as well as dilution and dispersion).

All available hydrological information were used to construct a CSM and numerical model of the CPP. A steady state, 5-layer numerical model, based on a previous groundwater model of the area, was constructed to characterize the long-term groundwater flow conditions at the site. The hydrostratigraphic units included in the model were the UCU, UA, and LCU. The DA and DCU were not included in the model. Calibration of the model focused on simulating mean groundwater elevations for 95 wells at the site by modifying hydraulic parameters for the different hydrostratigraphic units, alongside river and general head boundary conductance. The calibrated model represents a reasonable match to the observed head and sulfate concentration data. ritmate tor preferance constant marsport tems in the model. Boron, suitare, and D.
Inspects is likely to be affected by both chemical and physical attenuation mechanisms (sorption and/or precipitation reactions as well as

The calibrated model was used to predict the sulfate concentration for two closure scenarios using information provided in the Draft CCR Final Closure Plan (Golder Associates [Golder], 2022) including:

- **Scenario 1:** closure in place (CIP) including removal of CCR from the eastern portion of AP1, consolidation into the western portion of AP1, and construction of a cover system over the remaining CCR, and;
- **Scenario 2:** closure by removal (CBR) including removal of all CCR and regrading of the removal area.

Prior to the simulation of these scenarios, a dewatering simulation was included, which simulated the removal of free liquids from AP1 prior to the implementation of the two closure scenarios.

Differences exist in the timeframes to reach the GWPS for most monitoring wells between CIP and CBR. In general, the simulated groundwater concentrations in the monitoring wells within the UA will achieve the GWPS in 15 years for both the CIP and CBR closure scenarios, with the exception of well G301 in the CIP scenario. The predicted delayed reduction in concentration at well G301 is a result of the well being located along the flow path of the residual sulfate concentrations released into native geologic materials prior to closure. Reduced percolation rates through the consolidation area within AP1 in the CIP scenario means that the residual sulfate concentrations require a longer time period to migrate through native geologic materials.

Results of groundwater fate and transport modeling conservatively estimate that groundwater concentrations will attain the GWPS for all constituents identified as potential exceedances of the GWPS in the UA monitoring wells within 59 years of closure implementation for CIP and 15 years for CBR. The residual sulfate plumes from the calibrated model remain in close proximity to AP1 and have been simulated to decline below the GWPS (400 milligrams per [mg/L]) within 59 years for CBR. The residual plume in the CIP scenario will take longer in a small area at the northwest corner of AP1 due to the reduced infiltration rates below the cover system.

1. INTRODUCTION

1.1 Overview

In accordance with the requirements of Part 845 (IEPA, 2021), Ramboll has prepared this GMR on behalf of the CPP, operated by IPGC. This report will apply specifically to the CCR unit referred to as AP1 (**Figure 1-1**). However, information gathered to evaluate other CCR units at the CPP regarding geology, hydrogeology, and groundwater quality is included, where appropriate. AP1 is a 23-acre, unlined SI used to manage CCR and non-CCR waste streams at the CPP. Its total storage capacity is approximately 300 acre-feet. This GMR presents and evaluates the results of predictive groundwater modeling simulations for two proposed closure scenarios, including CCR consolidation and CIP, and CBR scenarios summarized below.

- **Scenario 1:** CIP including removal of CCR from the eastern portion of AP1, consolidation into the western portion of AP1, and construction of a cover system over the remaining CCR.
- **Scenario 2:** CBR including removal of all CCR and regrading of the removal area.

1.2 Previous Groundwater Modeling Reports

Several reports containing groundwater modeling have been completed at the CPP. The information presented in this GMR includes data collected in support of the previous groundwater models as well as data collected as part of a 2021 field investigation to support development of a HCR (Ramboll, 2021a). The HCR was provided as an attachment to the initial operating permit application required by 35 I.A.C. § 845.230. Previous groundwater modeling reports completed for the various CCR units located at the CPP include, but are not limited to, the following (recent to oldest): gerding geology, hydrogeology, and groundwater quality is included, where appropriate
garding geology, hydrogeology, and groundwater quality is included, where appropriately is approximately 300 acre-feet. This GMR present

• *Natural Resources Technology, Inc. (NRT), January 24, 2017. Hydrostatic Modeling Report. Coffeen Power Station, Coffeen, Illinois.*

Utilized the Hydrologic Evaluation of Landfill Performance (HELP) model to predict percolation from Ash Pond No. 2 (AP2) and evaluate AP2 hydrostatic conditions in response to the proposed cover system as described in the Revised 30% Closure Design Package.

• *NRT, January 24, 2017. Groundwater Modeling Report. Coffeen Power Station, Coffeen, Illinois.*

Included simulations of the site hydrology, the extent of CCR leachate impacts on groundwater, and the effect of pond closure on groundwater quality.

1.3 Site Location and Background

The CPP is located in Montgomery County, in central Illinois, within Section 11 Township 7 North and Range 7 East (**Figure 1-1**). The CPP is approximately two miles south of the city of Coffeen and about eight miles southeast of the city of Hillsboro, Illinois. AP1 is located between the two lobes of Coffeen Lake (identified as "Coffeen Lake" and "Unnamed Tributary" on **Figure 1-1** and **Figure 1-2**) to the west, east, and south, and is bordered by agricultural land to the north. The approximately 1,100-acre Coffeen Lake was built by damming the McDavid Branch of the East Fork of Shoal Creek in 1963 for use as an artificial cooling lake for the CPP. Historically, several coal mines were operated at depth in the vicinity of the CPP as well as the US Minerals processing facility located to the north. **Figure 1-2** is a site map showing the location of AP1 (Part 845 regulated CCR unit and subject of this GMR), AP2, Gypsum Management Facility Recycle Pond

(GMF RP), Gypsum Management Facility Gypsum Stack Pond (GMF GSP), and Landfill (LF). A surface water pond southwest of the LF collects overflow from the LF, this feature does not contain CCR. The area near AP1 will hereinafter be referred to as the Site.

1.4 Site History and CCR Units

The CPP was a coal-fired electrical generating plant that began operation in 1964. The plant initially burned bituminous coal from Illinois and CCR from the coal fired units was disposed of in AP1. AP2 was also utilized in the early 1970's and AP1 was reconstructed in 1978. Both of these units were used until the mid-1980's. Beginning in 2010, CCR material was placed in the LF and GMF units (*i.e.,* GMF RP and GMF GSP). All approximate dates of construction of each successive stage of the CCR units at the CPP are included in the groundwater model and described here.

AP1: This SI (also known as the Bottom Ash/Recycle Pond) is a reclaimed ash pond that was reconstructed utilizing the existing earthen berms with reinforcement, as provided by Water Pollution Control Permit 1978-EA-389 issued by the IEPA on May 26, 1978. AP1 (existing unlined SI) covers an area of approximately 23 acres, has berms up to 41 feet above the surrounding land surface, and a volume of 300 acre-feet. It primarily received bottom ash and low volume wastes from floor drains in the main power block building. Several years ago, air heater wash and boiler chemical cleaning wastes were directed to AP1, but this practice was discontinued. The bottom ash was periodically removed for beneficial uses by a third-party contractor. Sluicing of waste to AP1 ceased prior to November 4, 2019. 21. AP2 was also utilized in the early 1970's and AP1 was reconstructed in 1978. Both o
its were used until the mid-1980's. Beginning in 2010, CCR material was placed in the
MF units (*i.e.*, GMF RP and GMF GSP). All appro

AP2: AP2 is a closed (IEPA approved) SI with a surface area of approximately 60 acres and berms 47 feet higher than the surrounding land surface. AP2 was originally removed from service and capped in the mid 1980's. A clay and soil cap was placed on the surface of the pond with contouring and drainage provided to direct storm water to four engineered revetment down drain structures. Prior to capping, this pond was identified as Outfall 004 in the facility National Pollutant Discharge Elimination System (NPDES) operating permit, IL0000108. Additional closure activities include the construction of a geomembrane cover system that began in July 2019 and was completed on November 17, 2020. The construction was completed in accordance with the Closure and Post Closure Care Plan approved by the IEPA on January 30, 2018.

GMF GSP: The 77-acre GMF GSP received blowdown from the air emission scrubbers and was put into operation in 2010. Construction of the GMF GSP was in accordance with Water Pollution Control Permit 2008-EA-4661 and features a composite 60- one thousandth of an inch (mil) highdensity polyethylene (HDPE) liner with 3 feet of recompacted soil with a hydraulic conductivity of 1×10^{-7} centimeters per second (cm/s) with internal piping and drains to collect contact water. Construction of the unit required excavation to approximately 603 feet North American Vertical Datum of 1988 (NAVD88), removal of the sands and silts of the UA prior to construction of the liner, and installation of a groundwater underdrain system to eliminate inward pressure on the liner prior to placement of CCR. The GMF GSP underdrain was actively pumped during construction but is no longer actively pumped. IPGC ceased receipt of waste to the GMF GSP prior to April 11, 2021.

GMF RP: The 17-acre GMF RP received blowdown from the air emission scrubbers and was put into operation in 2010. Construction of the GMF RP was in accordance with Water Pollution Control Permit 2008-EA-4661 and features a composite 60-mil HDPE liner with 3 feet of recompacted soil with a hydraulic conductivity of 1×10^{-7} cm/s with internal piping and drains to collect contact water. Construction of the unit required excavation to approximately 601 feet

NAVD88, removal of the sands and silts of the UA prior to construction of the liner, and installation of a groundwater underdrain system to eliminate inward pressure on the liner prior to placement of CCR. The GMF RP underdrain is a passive, gravity drained system. IPGC ceased receipt of waste to the GMF RP prior to April 11, 2021.

LF: Fly ash was managed in a permitted composite lined landfill constructed in 2010. The LF has an active groundwater underdrain system that is currently being pumped. Additionally, the ash landfill leachate collection system is restricted by rule to no more than one foot of leachate on the composite liner. An IEPA groundwater monitoring program is in effect for the GMF GSP and GMF RP (under Bureau of Water), and LF (under Bureau of Land). nomini leachate collection system is restricted of y tile to no more to marke that the collection system is in effect for the GMF GSP and under Bureau of Water), and LF (under Bureau of Land).

2. SITE GEOLOGY AND HYDROGEOLOGY

2.1 Stratigraphy

The geology and hydrogeology of AP1 are described in detail in the HCR (Ramboll, 2021a) and summarized below.

The unlithified stratigraphy within and immediately surrounding AP1 consists of the following in descending order: fill material and CCR; clays and silts (Loess Unit); gravelly clay till and sandy materials, absent in some locations (Hagarstown Member); a weathered till zone and sandy, silt, or clay till (Vandalia Member); silt and sandy silt/clay unit (Mulberry Grove Member); silty clay diamicton (Smithboro Member); sand and sandy silt/clay, absent in some locations (Yarmouth Soil); and clay and silt with some sand (Lierle Clay Member). The unlithified units overlay Pennsylvanian‐age limestone, sandstone, and minor coal beds (Bond Formation). The Bond Formation bedrock was not encountered in any borings advanced at the CPP, so site-specific information is not available. e unimient so tadyaphy when and minetaledry solutioning wer a couster of the mathematic states and the capacity of the capacity of the particle of the capacity of the capacity of the capacitom Member); a weakhered till zon

CCR consisting of bottom ash and other non-CCR waste is present within AP1 at a thickness of up to 18 feet, as estimated from borings advanced within AP1, and an average thickness of 10 feet. However, CCR materials may be thicker near former drainage features in localized areas eroded through the loess and clay (Ramboll, 2021a). One such former drainage feature is located in the northeast corner of AP1 and ash fill may be in contact with the sandy portion of the Hagarstown Member similar to features observed at AP2. Non-CCR fill material consisting of silty clay, sandy lean clay, or lean clay with sand, with trace amounts of fine gravel comprises the berms surrounding AP1.

The Loess Unit is the uppermost unlithified unit identified at the CPP. This unit is comprised of the combined Roxana and Peoria Silt and extends from beneath the topsoil, derived from the loess, to the top of the Hagarstown Member. The loess has been classified as silt or clayey silt, with minor amounts of sand. The Loess Unit ranges in thickness from 0 feet (absent) to 16 feet, and was generally 8 to 14 feet thick, where present near AP1. The Loess Unit is generally considered unsaturated, and the UA is recharged by precipitation that percolates through this unit.

The Hagarstown Member (also referred to as Hagarstown Beds) exhibits two units: the first unit consisting of the gravelly clay till and the second consisting of sandy material overlying the Vandalia Member. The clay till portion had varying thicknesses ranging from approximately 2 to 6 feet as observed adjacent to AP1 (Ramboll, 2021a). The sandy portion of the Hagarstown, where present, was typically encountered between 9 and 34 feet below ground surface (bgs) near AP1, and is generally 1 to 5 feet thick, although thicknesses up to 7 feet have been observed north of the LF (Ramboll, 2021d; Ramboll, 2021e). The composition of the sandy portion of the Hagarstown unit varies across the CPP and has been classified as gravelly till, poorly sorted gravel, well sorted gravel, sand, and silty sand. Based on historic topography, the Hagarstown Member is not present in former drainage features present along the banks of Coffeen Lake and the Unnamed Tributary. During construction of the LF, GMF GSP, and the GMF RP, the Loess Unit and portions of the Hagarstown Member were excavated to facilitate construction.

The Vandalia (*i.e.,* till) Member is a sandy/silty till with thin, discontinuous lenses of silt, sand, and gravel. The Vandalia Member was encountered between 1.5 and 34 feet bgs in all borings

advanced at the CPP. The Vandalia Member typically ranged in thickness from 11.7 feet in the northern portion of the CPP, to 31.0 feet between the GMF GSP and the GMF RP. Similar to the observed top elevation of the Hagarstown Member, the top of the Vandalia Member declines in elevation near Coffeen Lake and topographic drainage features. This unit is relatively thick throughout the CPP, with an average thickness of over 15 feet (Hanson Professional Services, Inc. [Hanson], 2009).

The Mulberry Grove (*i.e.,* silt) Member typically consists of a thin, lenticular unit of gray sandy silt (Willman et al., 1975). It represents the interval between the retreat of the glacier that deposited the Smithboro Member and the advance of the glacier that deposited the Vandalia Member. At the CPP, the Mulberry Grove Member is represented by gray sandy silt layers deposited in depressions found in the surface of the underlying Smithboro Member. This unit was absent in many borings through the central portion of the CPP from south to north, and is generally less than 2 feet thick, but was measured at up to 4.9 feet thick near the GMF GSP (Hanson, 2009). is multiplery lorove (i.e., siit) member typically consists or a timi, ientotual runt or gray
en multiplery lorove (i.e., siit) member and the advance of the glader that deposited the Smithboro Member and the advance of th

The Smithboro (*i.e.,* till) Member is described as a gray, compact, silty, clayey diamicton that ranges in thickness from 6.7 to 21.2 feet northwest of the LF.

The Yarmouth Soil is described as the weathered zone on the Kansan drift, but in some places, it consists of accretionary deposits of fine sediment and organic material that accumulated in poorly drained areas on the surface of the Kansan deposits. Historical borings in the northern portion of the CPP which encountered the Yarmouth were summarized previously by Hanson (2009) as ranging in thickness from 0 feet (absent) to 5.1 feet.

The Lierle Clay Member is the uppermost member of the Kansan Stage Banner Formation. It is described as an accretion gley with clay, silt, and some sand. It was encountered by Hanson (2009) in all but a few borings on site. During the 2021 investigation, the top of the Lierle Clay was observed between 54 and 57 feet bgs. No borings advanced at the CPP penetrated the full thickness of the Banner Formation.

Pennsylvanian-age Bond Formation bedrock was not encountered in any borings advanced at the CPP, so site-specific information is not available.

2.2 Hydrogeology

Regionally, the water table conforms to the topographic features of the land surface. Recharge occurs in the uplands and flows towards drainage features. Moderate thicknesses of unconsolidated materials fill shallow valleys or are present on the uplands bordering the main valleys. These materials contain thin and discontinuous deposits of sand and gravel.

2.2.1 Groundwater Flow

Monitoring well locations are illustrated in **Figure 2-1**. Monitoring well locations and construction details are summarized in **Table 2-1**. Overall groundwater flow within the UA is divided towards the two lobes of Coffeen Lake. Groundwater generally flows from the center of the CPP west towards Coffeen Lake, and east towards the Unnamed Tributary, the eastern lobe of Coffeen Lake, and the discharge flume, resulting in a groundwater divide (high) running through the middle of the CPP (**Figure 2-2** and **Figure 2-3**). Groundwater flows north to northeast across AP1 toward the former discharge structure and Unnamed Tributary. Although elevations vary

seasonally, the groundwater flow direction in the UA is consistent and likely controlled by the proximity and hydraulic connection to Coffeen Lake.

2.2.2 Hydraulic Properties

Over 100 monitoring wells have been installed since 2006 to monitor groundwater conditions around the five CCR units at the CPP for both State and Federal groundwater compliance programs. Six hydrostratigraphic units were described in detail in the HCR (Ramboll, 2021a) and are summarized as follows:

- **CCR**: This unit is composed of CCR, consisting primarily of bottom ash. This also includes earthen fill deposits of predominantly silt and clay materials from on-site excavations that were used to construct berms and roads surrounding the various impoundments across the CPP. Laboratory testing of one CCR (ash) sample from AP1 had a vertical hydraulic conductivity of 8.8×10^{-5} cm/s.
- **UCU**: Consists of the Loess Unit and the upper clayey portion of the Hagarstown Member which has generally lower vertical permeability and generally greater than 60 percent fines (Ramboll, 2021a). This unit was encountered across most of the CPP, with the exception of the eastern edges of AP1 near the Unnamed Tributary where the unit was eroded following deposition or locations where it has been excavated for construction. Vertical hydraulic conductivities based on laboratory testing ranged from 1.3×10^{-8} to 5.0×10^{-7} cm/s.
- **UA** This unit consists primarily of sand and sandy silts and clays at the base of the Hagarstown Member and, in some locations, the uppermost weathered sandy clay portion of the Vandalia Member. This unit is absent in several locations due to weathering and in others due to excavation during construction of CCR Units. Field hydraulic conductivity tests indicated hydraulic conductivities ranged from 1.7×10^{-5} to 9.1×10^{-3} cm/s near AP1. Laboratory testing of one UA sample, collected near the GMF RP, had a vertical hydraulic conductivity of 1.6 x 10-4 cm/s (Ramboll, 2021a).
- **LCU**: This unit is composed of the sandy clay till of the Vandalia Member, the silt of the Mulberry Grove Formation, and the compacted clay till of the Smithboro Member. The unit underlies the UA and was encountered in all boring locations on the CPP. Results from laboratory tests completed for vertical hydraulic conductivity indicate the Vandalia Member has a very low vertical hydraulic conductivity. Field hydraulic conductivity tests indicated hydraulic conductivities from 4.0 x 10^{-8} to 3.4 x 10^{-5} cm/s; however, these likely reflect the isolated and discontinuous sandy lenses. Vertical hydraulic conductivities based on laboratory testing were from 1.3×10^{-8} to 5.0×10^{-7} cm/s. e summarized as follows:
 CCCR: This units is composed of CCR, consisting primarily of bottom ash. This also inclu
 erarthen fill deposits of predominantly silt and clay materials from on-site excavations

were used t
- **DA**: This unit consists primarily of sandy silt and sands of the Yarmouth Soil, which are thin (less than 5 feet) and discontinuous across the CPP. Field hydraulic conductivity tests indicated hydraulic conductivities from 8.7×10^{-5} to 1.7×10^{-3} cm/s within the DA.
- **DCU**: This unit underlies the DA and is composed of the Banner Formation, of which the thick Lierle Clay is the first encountered unit. No boring penetrated the full thickness of this formation.

2.2.3 Groundwater Elevation Data

During the 2021 Part 845 investigation, groundwater elevations in the UA ranged from approximately 591 to 625 feet NAVD88 across the CPP. Groundwater elevations were typically

highest towards the northern extent of the CPP, near the GMF GSP and GMF RP, except monitoring well G307 south of AP1, which consistently had the highest groundwater elevation. Groundwater elevations were lowest near the Unnamed Tributary and east of AP1 towards Coffeen Lake. Groundwater elevations in the vicinity of AP1 were typically from 591 to 621 feet NAVD88, with the exception of G307 as noted above, which was typically around 624 feet NAVD88 (**Figure 2-2** and **Figure 2-3**).

No seasonal variation has been observed in the UA monitoring wells, and any seasonal responses may be muted by the proximity and hydraulic connection to Coffeen Lake.

2.2.4 Mining Activity

Several coal mines, both strip and underground types, previously operated in Montgomery County, Illinois. Three mines - the Hillsboro Mine (Illinois State Geological Survey [ISGS] Mine No. 871), the Clover Leaf No. 4 Mine (ISGS Mine No. 442), and the Clover Leaf No. 1 Mine (ISGS Mine No. 3001) – were operated as room and pillar mines in the vicinity of the site beginning as early as 1889. The mines extracted coal from the Herrin (No. 6) Coal at depths of approximately 500 to 535 feet bgs (ISGS, 2019). All nearby mining operations ceased in 1983. Seasonal vanatorinats been observed in the UA monttoring weiss, and any seasonal vanisor and the UA monttoring the multiply and hydraulic connection to Coffeen Lake.

2.4 **Mining Activity**

Werel coal mines, both strip and

The Hillsboro Mine showed indications of small-scale faulting, roof stability issues, and floor heaving. Mine shafts, processing facilities, and some historic coal storage associated with these historic mines were located south of AP1. AP1 directly overlies the Hillsboro Mine. AP1 is outside of the buffer zone of the Clover Leaf No. 4 and Clover Leaf No. 1 mines (Ramboll, 2021a).

3. GROUNDWATER QUALITY

3.1 Groundwater Classification

Per 35 I.A.C. § 620.210, groundwater within the UA at AP1 meets the definition of Class I - Potable Resource Groundwater based on the following criteria:

- Groundwater in the UA is located 10 feet or more below the land surface and
- Within a geologic material which is capable of a hydraulic conductivity of 1×10^{-4} cm/s or greater using a slug test.

Field hydraulic conductivity tests performed in the UA near AP1 in 2021 had a geometric mean of 2.0 x 10^{-3} cm/s (Ramboll, 2021a). Based on this information, groundwater is classified as Class I – Potable Resource Groundwater.

However, background (upgradient) groundwater originates from areas southwest of AP1 that have historically been used for coal storage and present a potential alternate source for groundwater impacts.

3.2 Potential Groundwater Exceedances

A review and summary of data collected from 2015 through 2021 for parameters with GWPSs listed in 35 I.A.C. § 845.600 is provided in the HCR (Ramboll, 2021a). Concentration results presented in the HCR were compared directly to 35 I.A.C. § 845.600 GWPSs to determine potential exceedances. The results are considered potential exceedances because the results were compared directly to the standard and did not include an evaluation of background groundwater quality or utilize the statistical methodologies proposed in the GMP (Ramboll, 2021c) attached to the operating permit application. Within a geologic material which is capable of a hydraulic conductivity of 1×10^4 cm/s
greater using a slug test.
eled hydraulic conductivity tests performed in the UA near AP1 in 2021 had a geometric cloud
of a view

Groundwater concentrations from 2015 to 2021 are summarized in the History of Potential Exceedances (Ramboll, 2021b) (attached to the operating permit application) and are considered potential exceedances because the methodology used to determine them is proposed in the Statistical Analysis Plan (Appendix A to the GMP, Ramboll 2021c), which has not been reviewed or approved by IEPA at the time of submittal of the Part 845 operating permit application.

The History of Potential Exceedances attached to the operating permit application summarizes all potential groundwater exceedances following the proposed Statistical Analysis Plan. The following potential exceedances were identified:

- Boron determined at well G313.
- Cobalt determined at well G314.
- pH (lower limit) determined at well G312.
- Sulfate determined at wells G301, G303, G304/G307, G305, G307D, G308, G309, G310, G311, G312, G313, G314, G314D, G315, and G317.
- TDS determined at wells G303, G304/G307, G305, G307D, G308, G309, G310, G311, G312, G313, G314, G315, and G317.

A Technical Memorandum (**Attachment A**) was prepared by Geosyntec Consultants, Inc. (Geosyntec, 2022a), *Draft Evaluation of Potential Groundwater Protection Standard Exceedances, Coffeen Ash Pond No.1, Coffeen Illinois*, to further evaluate potential GWPS exceedances. The

results of the evaluation demonstrated that the potential GWPS exceedances of cobalt in well G314 and pH in well G312 are not related to AP1 based on several lines of evidence presented in the Technical Memorandum. Since potential GWPS exceedances for cobalt and pH are not related to AP1, these parameters will not be discussed further in this GMR.

RAFT

4. GROUNDWATER MODEL

4.1 Overview

Data collected at the Site from 2015 to the 2021 field investigation were used to update an existing groundwater model of the CPP (NRT, 2017b). The updated model was then used to evaluate the results of predictive groundwater modeling simulations for two proposed closure scenarios, including CCR consolidation and CIP, and CBR. The modeling results are summarized and evaluated in this GMR. The associated model files are included as **Appendix B**.

4.2 Description of Existing Model

The NRT (2017b) contaminant fate and transport model simulated boron and was performed to support closure of AP2 using MODFLOW and MT3DMS. AP1, GMF GSP, GMF RP, and LF were present within the previous model domain.

The NRT (2017b) modeling consisted of the following:

- Steady-state MODFLOW model was developed to represent site conditions for 2016. This model was calibrated to a set of groundwater elevation data collected during November 2016.
- The hydraulic properties from the steady-state model were used in the calibration of the transient MODFLOW and MT3DMS models which simulated groundwater flow and transport at the AP2 from 1970 to 2017. Boron concentrations collected in August 2016 were used to calibrate the transport model. enarios, including CCR consolidation and CIP, and CBR. The modeling results are summ

and evaluated in this GMR. The associated model files are included as **Appendix B.**

2 **Description of Existing Model**

NRT (2017b) cont
- Predictive simulations to estimate future boron concentrations for a baseline (no action) and capping closure scenario for AP2 were completed. Closure action was modeled over a period of 1,500 years, beginning in January 2018.
- Predicted boron concentrations were simulated to reach compliance for CIP at AP2 after 101 years (NRT, 2017b). These modeling results were part of the closure plan approved by IEPA on January 30, 2018.

4.3 Conceptual Model

The HCR (Ramboll, 2021a) forms the foundation of the AP1 hydrogeological setting. The AP1 overlies the recharge area for the underlying transmissive geologic media, which are composed of unlithified deposits.

4.3.1 Hydrogeology

As discussed in **Section 2.2**, groundwater flow in the UA at the CPP is divided towards the two lobes of Coffeen Lake. The loess of the UCU and sands of the UA are hydraulically connected. The groundwater flow in the silts and clays of the UCU and LCU are expected to be primarily vertical. The Hagarstown member is where the majority of the horizontal migration is expected to occur. The hydrogeological CSM consists of the following layers:

- Hagarstown Loess Unit (*i.e.,* UCU) Loess Unit and the upper clayey portion of the Hagarstown Member.
- Hagarstown Member (*i.e.,* UA) sand and sandy silts and clays at the base of the Hagarstown Member and, in some locations, the uppermost weathered sandy clay portion of the Vandalia Member.
- Vandalia Member/Mulberry Grove Member (*i.e.,* LCU) unweathered sandy clay till and discontinuous silts.
- Smithboro Till (*i.e.,* LCU) compacted clay till of the Smithboro Member.

The hydrostratigraphic units included in the model were the UCU, UA, and LCU. The DA and DCU were not included in the model, which includes consistency with the original model (NRT, 2017b). No potential GWPS exceedances have been observed in the DA. This, coupled with the limited groundwater data available for the DA and DCU, meant that these layers were not included in the model. Therefore, the Smithboro Till (*i.e.,* LCU) represents the lower boundary of the CSM.

Surfaces for each of the three major geological units (Loess Unit, Hagarstown Member, Vandalia/Mulberry Grove Member and Smithboro Till Member) were taken from the NRT model (2017b). The NRT model (2017b) used available information from well logs to interpolate the top and base of the UA.

4.3.2 Extent and Boundaries

The United States Geological Survey (USGS) National Map places the CPP within the East Fork Shoal Creek watershed subbasin (Hydrologic Unit Code [HUC] 071402030303).

The CPP CSM extent is bounded by a hydrological catchment (watershed) divide to the east based on watershed data from USGS. Along the north, south, and east, the model boundary has been placed along known waterbodies as much as possible. As such, it is assumed groundwater inflow from adjacent watersheds is negligible through both the UA and LCU.

The Coffeen Lake water levels are managed an average elevation 591.0 feet NAVD88. Coffeen Lake and Unnamed Tributary are the receiving body of water for surface water in the area encompassed by the CSM.

Infiltration of precipitation to the groundwater table is applied as recharge at the site. Groundwater in the UCU migrates downward into the Hagarstown Formation. As discussed in **Section 2.2.1**, the Hagarstown Formation is considered the UA for groundwater adjacent to AP1.

4.3.3 Ash Pond No. 1

AP1 is constructed such that the earthen berm and base are in contact with the UCU with exception of limited areas in the northeast of the SI where the UCU and UA have been eroded and the berm and base of CCR are in contact with the LCU. Findings from the HCR (Ramboll, 2021a) indicate that AP1 does influence the UA flow system, where there is a component of radial flow from AP1. However, this radial flow system appears to be centered around the southwest corner of AP1 resulting in a northerly and easterly component of groundwater flow within the UA. oundwater data vailable for the DA and DCU, meant that these layers were not included
Codel. Therefore, the Smithboro Till (i.e., LCU) represents the lower boundary of the CSN
unfectes for each of the three major geologica

Sulfate was selected for transport modeling. Sulfate is commonly used as an indicator parameter for contaminant transport modeling for CCR because: (i) it is commonly present in coal ash leachate; and (ii) it is mobile and typically not very reactive but conservative (*i.e.*, low rates of sorption or degradation) in groundwater.

4.5 Model Approach

4.5.1 Potential Groundwater Exceedances

A comparison of observed TDS concentrations to sulfate (**Figure A** below) indicates a statistically significant correlation between these parameters in UA wells where these potential exceedances were observed. Observed concentrations were transformed into Log10 concentrations for evaluation. The correlation coefficient (R^2) and p values (indicator of statistical significance) are also provided on Figure A. Higher R² values (*i.e.*, closer to 1) indicate stronger correlation between parameters. A correlation is considered statistically significant when the p value is lower than 0.05. The correlation has a p value less than the target of 0.05, indicating the correlation is statistically significant. The statistically significant correlation associated with sulfate concentrations indicate sulfate is an acceptable surrogate for TDS in the groundwater model, and concentrations of this parameter are expected to change along with model predicted sulfate concentrations.

Figure A. Sulfate Correlation with TDS in UA Wells

A potential exceedance of boron was also observed at one monitoring well, G313, in the vicinity of AP1, based on the History of Potential Exceedances (Ramboll, 2021b). Correlations between sulfate and boron for the same AP1 UA wells did not indicate a statistically significant correlation between these constituents. However, UA monitoring well G313 has potential exceedances of both sulfate and TDS along with the potential exceedance of boron (**Section 3.2**). Boron, like sulfate, is a common indicator parameter used for contaminant transport modeling of CCR; and boron is less likely than other constituents to be present in background groundwater from natural or other anthropogenic sources. The only significant source of boron is AP1. With potential

exceedances of boron, sulfate, and TDS present in the same well (G313) and having the same source (AP1), boron concentrations are expected to change along with model predicted sulfate concentrations.

4.5.2 Summary of Modeling Activities

A three-dimensional groundwater flow model was calibrated to represent the conceptual flow system described above. Prediction simulations were performed to evaluate the effects of closure (source control) measures (CCR consolidation and CIP and CBR scenarios) for the CCR units on groundwater quality following initial corrective action measures, which includes removal of free liquids (dewatering). **Figure 4-1** illustrates the calibration and predictive modeling timelines.

Three model codes were used to simulate groundwater flow and contaminant transport:

- Groundwater flow was modeled in three dimensions using MODFLOW 2005.
- Contaminant transport was modeled in three dimensions using MT3DMS.
- Percolation (recharge) was modeled using the results of HELP model.

Modeling steps are a summarized below:

- A steady state model was created in MODFLOW 2005 and used to simulate the mean groundwater flow conditions at the site. The model was calibrated to match mean groundwater elevations observed between 2015 to 2021 (**Table 4-1**).
- Transient flow models based off of the calibrated steady state model were used to simulate groundwater flow and transport for 42 years using MODFLOW 2005 and MT3DMS to simulate changes in site conditions through time and match currently observed concentrations of sulfate in groundwater (**Table 4-1**). our econtrol) measures (CCR consolidation and CIP and CBR scenarios) for the CCR unreading production and predictive action measures, which includes removal counting (devatering). **Figure 4-1** illustrates the calibration a
- Prediction simulations began with a 2-year dewatering period simulated in MODFLOW 2005 and MT3DMS where heads were reduced within the CCR unit and concentrations were removed from CCR removal areas.
- Prediction simulations resumed for CIP and CBR following the 2-year dewatering period using the results of HELP modeling as input values for recharge rates in the construction areas.
- The prediction simulations were run using MODFLOW 2005 and MT3DMS to estimate the time for sulfate concentrations to meet the GWPS in the compliance wells and to evaluate the differences between the two closure scenarios.

5. MODEL SETUP AND CALIBRATION

5.1 Model Descriptions

For the construction and calibration of the numerical groundwater flow model for the site, Ramboll selected the model code MODFLOW, a publicly-available groundwater flow simulation program developed by the USGS (McDonald and Harbaugh, 1988). MODFLOW is thoroughly documented, widely used by consultants, government agencies and researchers, and is consistently accepted in regulatory and litigation proceedings. MODFLOW uses a finite difference approximation to solve a three-dimensional head distribution in a transient, multi-layer, heterogeneous, anisotropic, variable-gradient, variable-thickness, confined or unconfined flow system—given user-supplied inputs of hydraulic conductivity, aquifer/layer thickness, recharge, wells, and boundary conditions. The program also calculates water balance at wells, rivers, and drains.

MODFLOW was developed by USGS (McDonald and Harbaugh, 1988) and has been updated several times since. Major assumptions of the code are: (i) groundwater flow is governed by Darcy's law; (ii) the formation behaves as a continuous porous medium; (iii) flow is not affected by chemical, temperature, or density gradients; and (iv) hydraulic properties are constant within a grid cell. Other assumptions concerning the finite difference equation can be found in McDonald and Harbaugh (1988). MODFLOW 2005 was used for these simulations with Groundwater Vistas 7 software for model pre- and post- processing tasks (Environmental Simulations, Inc., 2017). orinented, widely used by consultants, government agencies and researchers, and is
nosistently accepted in regulatory and litigation proceedings. MODFLOW uses a finite differ
proximation to solve a three-dimensional head d

MT3DMS (Zheng and Wang, 1998) is an update of MT3D. It calculates concentration distribution for a single dissolved solute as a function of time and space. Concentration is distributed over a three-dimensional, non-uniform, transient flow field. Solute mass may be input at discrete points (wells, drains, river nodes, constant head cells), or distributed evenly or unevenly over the land surface (recharge).

MT3DMS accounts for advection, dispersion, diffusion, first-order decay, and sorption. Sorption can be calculated using linear, Freundlich, or Langmuir isotherms. First-order decay terms may be differentiated for the adsorbed and dissolved phases.

The program uses the standard finite difference method, the particle-tracking-based Eulerian-Lagrangian methods and the higher-order finite-volume total-variation-diminishing (TVD) method for the solution schemes. The finite difference solution has numerical dispersion for lowdispersivity transport scenarios but conserves good mass balance. The particle-tracking method avoids numerical dispersion but was not accurate in conserving mass. The TVD solution is not subject to significant numerical distribution and adequately conserves mass, but is numerically intensive, particularly for long-term models such as developed for this model. The finite difference solution was used for this simulation.

Major assumptions of MT3DMS are: (i) changes in the concentration field do not affect the flow field; (ii) changes in the concentration of one solute do not affect the concentration of another solute; (iii) chemical and hydraulic properties are constant within a grid cell; and (iv) sorption is instantaneous and fully reversible, while decay is not reversible.

The HELP model was developed by the United States Environmental Protection Agency (USEPA). HELP is a one-dimensional hydrologic model of water movement across, into, through, and out of

a landfill or soil column based on precipitation, evapotranspiration, runoff, and the geometry and hydrogeologic properties of a layered soil and waste profile. For this modeling, results of the HELP model, HELP Version 4.0 (Tolaymat and Krause, 2020) completed for the groundwater model were used to estimate the hydraulic flux from closure construction.

5.2 Flow and Transport Model Setup

The 2017 flow and transport models were retained and revised as appropriate to perform simulations for the AP1.

The modeled area was approximately 10,000 feet by 15,025 feet $(150,250,000$ square feet $[ft^2]$) centered on the CPP (**Figure 5-1**). The model boundaries along the northern and eastern edges of the model were selected to maintain sufficient distance from the CPP to reduce boundary interference with model calculations, while not extending too far past the extent of available calibration data. The eastern edge of the model also approximates topographic highs, surface water divides, and watershed boundaries.

The steady state MODFLOW model was calibrated to mean groundwater elevation collected from 2015 to 2021 as presented in **Table 4-1**. MT3DMS was run on the transient flow model and model-simulated concentrations were calibrated to observed sulfate concentration values at the monitoring wells from January 2015 to July 2021 as presented in **Table 4-1**. Multiple iterations of MODFLOW and MT3DMS calibration were performed to achieve an acceptable match to observed flow and transport data. For AP1, the calibrated flow and transport models were used in predictive modeling to evaluate the CIP and CBR closure scenarios. Prior to simulation of CIP and CBR, a dewatering phase, which simulated the removal of free liquid from the CCR material in the AP1 was completed. Closure scenarios were simulated by removing saturated ash cells from removal areas and using HELP modeled recharge values to simulate changes proposed in the closure scenarios. mulations for the AP1.

Ne modeled area was approximately 10,000 feet by 15,025 feet (150,250,000 square feet the note CPP (Figure 5-1). The model boundaries along the northern and eastem

the model were selected to mainta

5.2.1 Grid and Boundary Conditions

A five-layer, 326 x 211 node grid was established with a variable grid spacing between 25 and 100 feet (**Figures 5-2 through 5-6**), with a total number of 284,575 active cells.

The main body of Coffeen Lake is immediately adjacent to CPP on the west and south and the Unnamed Tributary borders CPP to the east. These surface water features form the southern, eastern, and western boundaries of the model. The northern boundary of the model domain is a general head boundary. Vertically, the model domain extends from the top of the saturated zone to the base of the Smithboro Member. The thick clays of the Banner Formation are relatively impermeable compared to the overlying unconsolidated sediments and provides a base for the model.

The northern boundaries for layers 3, 4, and 5 are general head boundaries placed to simulate flow in the sandier soils of the Hagarstown Member, Vandalia/Mulberry Grove Member and Smithboro Till composing the UA (layer 3), and LCU (layer 4 and 5). The northern boundary represents the regional flow conditions within these units. The eastern edge is no-flow boundary in all model layers.

Coffeen Lake is represented as a constant head boundary based on an average surface water elevation of 591.0 feet NAVD88. The constant head boundary was simulated with an elevation equal to 591.0 feet. The lake is in hydraulic connection with multiple layers within the model.

The bottom of the model was also a no-flow boundary. The top of the model was a timedependent specified flux boundary, with specified flux rates equal to the recharge rate. A specified mass flux boundary was used to simulate downward percolation of solute mass from the AP1. This boundary condition assigns a specified concentration to recharge water entering the cells within AP1, and the resulting concentration in the AP1 cells is a function of the relative rate and concentration of recharge water (water percolating from the impoundment) compared to the rate and concentration of other water entering the node.

5.2.2 Flow Model Input Values and Sensitivity

Evaluation of monitoring well data for the CPP has not identified statistically significant seasonal trends in groundwater flow or quality which could affect model applicability for prediction of transport. The MODFLOW model was calibrated to mean groundwater elevations from 2015 to 2021. Multiple iterations of MODFLOW calibration were performed to achieve an acceptable match to observed flow data.

Sensitivity analysis was conducted by changing input values and observing changes in the sum of squared residuals (SSR). Horizontal conductivity, vertical conductivity, and river and general head conductance terms were all varied by one order of magnitude (*i.e.*, between one-tenth and ten times) of the calibrated values. Recharge terms were varied between one-half and two times calibrated values. River stage was obtained from the 10 meter (m) Digital Elevation Model (DEM) from the United States Department of Agriculture/Natural Resources Conservation Service (USDA/NRCS) National Geospatial Center of Excellence (USDA/NRCS, 2022). The vertical error of the 10 m DEM is 0.82 m (2.7 feet); therefore, the stream stage was varied by adding and subtracting 2.7 feet. Where appropriate, drain stage was modified based on the DEM error. Where this was inappropriate, drain stage increased and decreased by 2 feet. General head boundary head terms were varied between 90 and 110 percent of calibrated values. The HFB was varied by increasing the hydraulic conductivity by a factor of 100 and 1,000. When the calibrated model was tested, the SSR was 351. Sensitivity test results were categorized into negligible, low, moderate, moderately high, and high sensitivity based on the change in the SSR as summarized in the notes in **Table 5-1**. **2.2. Flow Model Input Values and Sensitivity**

2.2. **Elow Model Input Values and Sensitivity**

valuation of monitoring well data for the CPP has not identified statistically significant sec

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5.2.2.1 Layer Top/Bottom

The top of the saturated zone was used as the top of the model. The elevations for the base of each hydrostratigraphic layer were obtained from the NRT model (2017b) and were imported as grid data into MODFLOW. The upper Loess Unit of the Hagarstown Member (UCU) was divided into two layers to accommodate the explicit inclusion of the CCR in AP1 and AP2. The sand and silts of the Hagarstown Member which form the UA were represented using a single layer. The LCU was represented by two layers, the upper LCU (layer 4) represents the unweathered Vandalia/Mulberry Grove Member and the lower LCU (layer 5) represents the Smithboro Member.

The UCU layer was split into two layers (layers 1 and 2) to simulate the construction of AP1 and AP2. Within AP1 and AP2, layer 1 represents ash fill and layer 2 represents the UCU present below the ash and above the UA. Outside of AP1 and AP2, both layers 1 and 2 represent the UCU. Layer 3 represents the UA and the LCU is present in layers 4 and 5. **Figures 5-7 through 5-11** show the bottom elevations of the five model layers. The resulting model layers represent the distribution and change in thickness of each water-bearing unit across the model domain. **Table A** below provides elevation and thickness information for the model layers and hydrostratigraphic units used in the model.

Table A. Flow Model Layer Descriptions

Notes:

¹ Elevation is measured in feet, referenced to NAVD88.

5.2.2.2 Hydraulic Conductivity

Hydraulic conductivity values and sensitivity results are summarized in **Table 5-1.** The spatial distribution of the hydraulic conductivities within the UCU, UA and LCU were considered homogenous. **Figures 5-12 through 5-16** show the spatial distribution of the hydraulic conductivity zones, AP1 and other units on site for each of the five model layers. Construction of the GMF units removed the sands and silts of the UA prior to construction of the liner, therefore the UA is absent beneath these units and liner hydraulic properties are assigned. Conductivity zones that did not have representative site data (*i.e.,* zones 19 and 21, representing the cells above the river cells and the disturbed sediments between the LF and GMF GSP, respectively) were determined through model calibration. and CCR

UA Hagarstown Member

(604.0-614.15) (580.0-612.0) (2.1

LCU Vandalia/Mulberry 600.9

Groue Member

(590.0-612.0) (578.0-594.0) (2.1

LCU Base of Coffeen Lake (578.0-594.0) (2.1

Network Member

(598.0-612.0) (578

Where available, hydraulic conductivity values were derived from field measured or laboratory tested values reported in the HCR (Ramboll, 2021a) (**Section 2.2.2**). No horizontal anisotropy was assumed. Vertical anisotropy was applied to conductivity zones to simulate preferential flow in the horizontal direction in these materials, and are presented as anisotropy ratio (Kh/Kv) in **Table 5-1**.

The model was highly sensitive to changes in horizontal and vertical hydraulic conductivity in zones 1 (UCU), 2 (UA), and 3 (LCU - unweathered Vandalia), and moderately sensitive to changes in horizontal and vertical hydraulic conductivity in zones 10 (CCR fill-AP1) and 19 (UCU-fill). The model exhibited a negligible to low sensitivity in the remaining zones for both horizontal and vertical conductivity.

5.2.2.3 Recharge

Recharge rates were determined through calibration of the model to observed groundwater elevations. For the calibration model, recharge was applied to the uppermost active layer and the rates varied based on different units, namely the AP1, AP2, GMF GSP, GMF RP, LF, Surface Water Pond, and Cooling Pond. Model inputs are summarized in **Table 5-1**. The distribution of recharge is shown in **Figure 5-17**. Changes in operational history, such as the addition of AP1 to the site in 1977 and the GMF units in 2010 as illustrated in **Figures 5-18 through 5-21,** have been

incorporated into the transient model simulation (**Table 5-2**). See **Section 5.2.3.1** for additional discussion of time discretization.

The model had a high sensitivity to changes in recharge in zones 1 (UCU) and 7 (CCR fill - AP1). The model had negligible to low sensitivity to changes in recharge in the remaining zones, with the exception of zone 6 (CCR fill - AP2), where the sensitivity was moderate.

5.2.2.1 Storage and Specific Yield

The flow calibration model did not use these terms because it was run at steady state. For the transport model, which was run as a transient simulation, no field data defining these terms were available so published values were used consistent with Fetter (1988). Specific yield was set to equal effective porosity values described in **Section 5.2.3.5**. The spatial distribution of the storage and specific yield zones were consistent with those of the hydraulic conductivity zones. The sensitivity of these parameters was tested by evaluating their effect on the transport model as described in **Section 5.2.3.6**.

5.2.2.2 River Parameters

Five river reaches were included in the model as head dependent flux boundaries that required inputs for elevation of the surface water, bottom of the stream, width, bed thickness, and bed hydraulic conductivity (**Table 5-1**). The five river reaches were the Unnamed Tributary east of the CPP (reach 0 and reach 5), the Unnamed Tributary west of the CPP (reach 1), ponded surface water west of the LF (reach 2), and the condenser cooling water discharge flume (reach 3). The river and drain information is summarized in **Table B** below.

The flow calibration model did not use these terms because it was run at steady state. Fo					
transport model, which was run as a transient simulation, no field data defining these ter available so published values were used consistent with Fetter (1988). Specific yield was equal effective porosity values described in Section 5.2.3.5. The spatial distribution of th storage and specific yield zones were consistent with those of the hydraulic conductivity z The sensitivity of these parameters was tested by evaluating their effect on the transport as described in Section 5.2.3.6.					
River Parameters 5.2.2.2					
the CPP (reach 0 and reach 5), the Unnamed Tributary west of the CPP (reach 1), ponded water west of the LF (reach 2), and the condenser cooling water discharge flume (reach 3 river and drain information is summarized in Table B below. Table B. River and Drain Information					
Name	Boundary Type	Length (feet)	Slope (ft/ft)		
Unnamed Tributary East	River	8959.0	-0.0031		
Unnamed Tributary East - downstream reach	River	1438.3	-0.0026		
Unnamed Tributary West	River	3436.5	-0.0098		
Ponded Surface Water West	River				
Condenser Cooling Flume	River				
Active Landfill Underdrain	Drain	2147.0			
Gravity Drain Recycle Pond	Drain	2181.8			

Table B. River and Drain Information

Notes:

ft/ft = feet per foot

In the absence of river geometry information, the DEM was used to estimate stream stage at the upstream and downstream limits of the Unnamed Tributary east of the CPP and the Unnamed Tributary west of the CPP. The surface water stages for the ponded surface water west of the LF and the Condenser Cooling Flume were constant (not sloped) and were also obtained from the DEM. For both Unnamed Tributaries (east and west), the slope of the river was then linearly interpolated along the reaches, providing an estimation of stream stage along the length of each reach for each model grid cell though which the river flows. Bed thickness was set at 2 foot and river width was set at 10 feet. The river bottom is set 3 feet below the stage for both the

Unnamed Tributaries. The downstream reach (reach 5) of the Unnamed Tributary is located in layer 5 of the model adjacent to the SI unit AP2, this layer represents the LCU-Smithboro till and has a low hydraulic conductivity. To increase connectivity of the tributary to the overlying layers, the hydraulic conductivity of the streambed was modified during calibration.

The Condenser Cooling flume stage is maintained at 604.0 feet and the ponded surface water west of the LF was maintained at 617.5 feet, and bed thicknesses for these reaches were set to 1 foot. The width of the Cooling Flume (approximately 52 feet) and ponded surface water west of the LF are larger than the grid cell dimensions (25 feet by 25 feet); therefore, the conductance term for both were based on the area of the cells which coincide with the flume and ponded water.

The model had low to moderate sensitivity to changes in river stage. The model had low to moderately high sensitivity to changes in river conductance, with the exceptions of reach 0 (Unnamed Tributary East) and reach 3 (Condenser Cooling Flume) which had high sensitivity.

5.2.2.3 Drain Parameters

The LF has an active underdrain, which is actively pumped to prevent more than 1-foot of groundwater head above the liner. This was estimated to be 603.5 feet. The GMF RP has a passive drain beneath the liner which discharges water towards the Unnamed Tributary east of the unit. This was estimated to be 600.5 feet. Both the active LF drain and passive GMF RP drain were placed in layer 4 (LCU) below the low hydraulic conductivity zones which represent the base of the lined units. A surface water drain in the north of the model was also included; the placement of this northern drain was determined using google earth imagery. The Northern drain appears to be a man-made feature and no hydrological data are available as to its flow conditions. Therefore, its implementation in the model as a drain makes the fewest assumptions of its interaction with the aquifer. This surface water drain is located in layer 1 and has an elevation of 622.0 feet. or. Ine watch or the Cooling Fillmen (approximately 2/ reet pand ponded surred water
or. In even watch or the Coling Fillment approximately 2/ reet pand ponded surred water
m for both were based on the area of the cells wh

The model had low sensitivity to changes in drain stage. The model had negligible to moderate sensitivity to changes in drain conductance, with the exception of reach 0 (Active LF Underdrain) where the model had moderately high sensitivity to changes in drain conductance.

5.2.2.4 GMF Unit Parameters

All GMF units (GMF GSP, GMF RP, and LF) have a similar liner construction (**Table C** below); they were all implemented into the model using horizontal flow barrier (HFB) package to represent the liner system on the sides of the units. The bottom of the liner is implemented by assigning the liner system hydraulic conductance to model layer 3 within the footprint of the pond. The base elevation of layer 3 within the footprint of the GMF units simulates the base elevation of the liner. The thickness of model layer 3 within the footprint of the pond was set to three feet. Removal of the sands and silts below the GMF units (as described in **Sections 1.4** and **2.1**) means that the liner is in direct contact with the Vandalia Member. The groundwater flow dynamics beneath/around the Ash Landfill and GMF Units is affected by several factors, including: removal of the Hagarstown Member from beneath the Units; presence of the construction dewatering systems around the units; and the lateral variability of lithology within the Hagarstown Member (Hanson, 2016). Drains discussed above were used to represent the underdrains associated with the GMF units. The hydraulic properties within the GMF units were set to represent the CCR.

Estimates of the hydraulic properties of each of the components within the liner system were derived using values from the HELP model; see **Section 5-1** for more information about HELP. For flow perpendicular to the layer orientation, as is the case in the liner where the hydraulic gradient is vertical for the base and horizontal for the sides of the pond, the harmonic mean was used to obtain the effective hydraulic conductivity (K_{eff}) (Fetter, 1988). The harmonic mean was determined by:

$$
K_{eff} = \frac{\sum b}{\sum \frac{b}{K}}
$$

where b is the layer thickness and K is the horizontal hydraulic conductivity.

HFB input parameters are presented in **Table 5-1**. The model had low to moderate sensitivity to changes in the hydraulic conductivity in the HFB.

Liner Component	Thickness (feet)	Hydraulic Conductivity (cm/s)	Hydraulic Conductivity $({\rm ft}/{\rm d})$
HDPE geomembrane (60 mil)	0.06	2.0×10^{-13}	5.7 \times 10 ⁻¹⁰
Recompacted Soil	3.0	1.0×10^{-7}	2.8×10^{-4}
Vertical Harmonic Mean of liner system	NA	NA	2.89×10^{-8}

Table C. Liner System Properties From Top to Bottom for the GMF GSP, GMF RP, and LF

* Estimated based on available information $ft/d = feet$ per day

 $NA = not applicable$

5.2.2.5 General Head Boundary

General head boundary conditions (GHB) were used along the northern boundary of the model for layer 3 through 5 (**Figures 5-4 through 5-6**). The GHB at the northern limit of the model represents groundwater entering the model domain from upgradient areas. The GHB is present in layers 3 through 5 and was used to simulate groundwater flow into the model via the UA and LCU. The groundwater levels used for the northern boundary of the model in layers 3 through 5 were estimated using the Dupuit equation for steady state flow in an unconfined aquifer with recharge. $K_{eff} = \frac{E}{\sum_{K}^{2} D}$

There b is the layer thickness and K is the horizontal hydraulic conductivity.

FB input parameters are presented in **Table 5-1**. The model had low to moderate sensitianges in the hydraulic conduct

The DEM of the site provided estimates of the surface water levels for Coffeen Lake on the west of the model (591 feet), and Rocky Ford Sportsman Club North Lake (604 feet) on the east of the model domain (refer to **Figure 5-1**). The calibrated ambient recharge to the UCU was used in the calculation of the groundwater level distribution at the northern boundary. The hydraulic conductivity value used in the Dupuit equation was estimated during model calibration.

This GHB was only applied to cells along the northern boundary where the base of the cell was below the calculated groundwater head for a given distance from the constant head boundaries, the head was determined by the Dupuit equation. Cell conductance was then calculated using the cells' saturated thickness and cell width, and hydraulic conductivity were based on cell hydraulic conductivities and adjusted if appropriate during calibration.

The GHB elevation for northern boundary in the UA was established during calibration (**Table 5-1**). The distance to the GHB head was set to 1, and the GHB conductivity was calculated using the cell width, cell thickness, and calibrated hydraulic conductivity from the model.

The sensitivity to changes in specified head was low to moderate, with the exception of reach 3 (Northern Model Boundary in LCU Layer 4) where the model sensitivity was high. The flow calibration model had a negligible sensitivity to changes in conductance.

5.2.3 Transport Model

MT3DMS input values are listed in **Table 5-2** and described below. Sensitivity of the transport model is summarized in **Table 5-3**.

Groundwater transport was calibrated to groundwater sulfate concentration ranges at each well as measured from the monitoring wells between 2015 (where available) and 2021. The transport model calibration targets are summarized in **Table 4-1**.

Sensitivity analysis was conducted by changing input values and observing percent change in sulfate concentration at each well from the calibrated model sulfate concentration. Effective porosity was varied by decreasing and increasing calibrated model values by 0.05. Storage values were multiplied and divided by a factor of 10, and specific yield by a factor of 2. The dispersivity values in the calibrated model were increased by a factor of 5 and 10. The sensitivity of the transport model to changes in the liner conductance was also investigated by increasing and decreasing the hydraulic conductivity of the liner by one order of magnitude (*i.e.*, between one-tenth and ten times). T3DMS input values are listed in **Table 5-2** and described below. Sensitivity of the trandol is summarized in **Table 5-3**.
T3DMS input values are listed in **Table 5-3**.

coundwater transport was calibrated to groundwater s

The transport model had a negligible to moderate sensitivity to changes in storage and specific yield (**Table 5-3**) as discussed in **Section 5.2.3.6**. The transport model ranged from negligible to moderate sensitivity to effective porosity and dispersivity as discussed in **Sections 5.2.3.5** and **5.2.3.7**, respectively. The sensitivity to the liner conductivity was negligible to low as discussed in **Section 5.2.3.2**.

5.2.3.1 Time Discretization and Stress Periods

The evolution of the CPP required changes to the hydraulic properties within the model; this is not possible in a single model where hydraulic properties as assumed to remain constant. As a result, the changes in the site (*e.g.,* inclusion of the GMF units) are simulated in three consecutive numerical models, as summarized in **Table D** below. The simulation length was revised from the existing model to extend to the current time (2022).

Notes: TR = transient model

5.2.3.2 GMF Units

Groundwater chemistry data from wells G215 (located adjacent to the GMF GSP), and wells G275 and G279 (located adjacent to the GMF RP), indicate an increase in sulfate concentrations post 2018 when compared with sulfate concentrations in adjacent wells. Sulfate concentrations in G215 have experienced further increases since 2021. Sulfate concentrations around the GMF RP tend to be higher than those around the GMF GSP, with elevated sulfate concentrations observed since 2015 (the earliest sampling date). Elevated sulfate concentrations along the southern boundary of the GMF RP are associated with historic groundwater impacts from AP2. However, wells G275 and G279 are located along the eastern boundary of the pond and have elevated sulfate concentrations. To simulate observed sulfate concentrations at these isolated wells (GMF GSP well G215, and GMF RP wells G275 and G279), the hydraulic conductivity of the liner (simulated using HFB) was increased to allow sulfate migration from the CCR unit in the transient model TR-3, as shown in **Figure B** below and **Table D** above.

Figure B. Liner Modification Zones

As part of the transport calibration process, the hydraulic conductivity of HFB reaches 11, 16, and 21 were modified to simulate the observed rises in sulfate. The changes are summarized in **Table 5-2**. Model sensitivity near the GMF ponds is discussed in the Draft Groundwater Modeling

Report, *GMF Gypsum Stack Pond and GMF Recycle Pond, Coffeen Power Plant, Coffeen Illinois* (Ramboll, 2022).

The monitoring wells associated with AP1 show negligible to low sensitivity to changes in the GMF liner conductivity (**Table 5-3**). AP1 Is located approximately 2,500 foot south of the GMF SIs, any changes in groundwater flow and transport will be minimal in proximity to AP1.

5.2.3.3 Initial Concentration

No initial concentrations were placed in the steady state flow calibration model. The flow model was run as transient and concentration was added to the model through recharge starting at the same time as the transient flow simulation. Modeling was performed for a sufficient period (42 years) to allow modeled concentrations in the primary transport layer (*i.e.,* UA) to reach recently observed levels.

Modeling was performed over three numerical models which mirror the operational developments at the CPP. **Table 5-2** provides an overview of how the source concentrations and recharge rates change through time.

5.2.3.4 Source Concentration

Five sources in the form of vertical percolation (recharge) and constant concentration cells were simulated in the CCR material for calibration (**Table 5-2**) (in chronological order): (i) percolation through CCR in AP2 (1970-2022), (ii) percolation through CCR in AP1 (1978-2022), (iii) percolation through CCR in GMF RP (2010-2022), (iv) percolation through CCR in GMF GSP (2010-2022), and (v) percolation through CCR in GMF LF (2010-2022). All five sources were simulated by assigning concentration to the recharge input. The CCR sources were also simulated with constant concentration cells placed where CCR was present (**Figures 5-18 through 5-21**) to simulate saturated CCR conditions. From the model perspective, this means that when the simulated water level is above the base of these cells, water that passes through the cell will take on the assigned concentration. All source concentrations were calibrated in the transport model to the sulfate concentration data collected from November 2015 to August 2021. The source concentrations applied to the recharge zones and saturated ash cells immediately below the recharge zones have the same concentration values. **Table 4-1** indicates that the background sulfate concentrations (identified with a "B" for background in the "CCR unit" column) at CPP show considerable variability across the site, from 11 mg/L (G286) to 770.0 mg/L (G288). No background sulfate concentration was applied to recharge beyond the source areas in the model. is initial concentrations were placed in the steady state flow calibration model. The flow
initial concentrations were placed in the steady state flow calibration model. The flow
member as the transient flow simulation. Mo

Because these are the sources of concentration in the model, the model will be highly sensitive to changes in the input values. For that reason, sensitivity testing was not completed for the source values.

5.2.3.5 Effective Porosity

Effective porosity for each modeled hydraulic conductivity zones were based on the NRT model (2017b), data from the HCR (Ramboll, 2021a), and literature values (Fetter, 2001) and are presented in **Table 5-2**.

The model had a negligible to moderate sensitivity to changes in porosity values (**Table 5-3**). The greatest sensitivity for porosity was moderate for the high porosity sensitivity test at

monitoring locations G305, G306, and G317. Moderate sensitivity at monitoring well G317 was also observed for the low porosity sensitivity test.

5.2.3.6 Storage and Specific Yield

The transport model had a negligible to low sensitivity to changes in storage and specific yield, with the exception of sensitivity at monitoring wells G306, G307, and G317, where sensitivity was moderate (**Table 5-3**).

5.2.3.7 Dispersivity and Diffusion

Physical attenuation (dilution and dispersion) of contaminants is simulated in MT3DMS. Dispersion in porous media refers to the spreading of contaminants over a greater region than would be predicted solely from the average groundwater velocity vectors (Anderson, 1979; Anderson, 1984). Dispersion is caused by both mechanical dispersion, a result of deviations of actual velocity at a microscale from the average groundwater velocity, and molecular diffusion driven by concentration gradients. Molecular diffusion is generally secondary and negligible compared to the effects of mechanical dispersion and only becomes important when groundwater velocity is very low. The sum of mechanical dispersion and molecular diffusion is termed hydrodynamic dispersion, or simply dispersion (Zheng and Wang, 1998). 2.3.7 **Dispersivity and Diffusion**

2.3.7 **Dispersivity and Diffusion**

sysical attenuation (dilution and dispersion) of contaminants is simulated in MT3DMS.

spersion in porous media refers to the spreading of contaminant

Longitudinal dispersivity was 10 feet in the UA and 1 foot in the UCU and LCU, with transverse and vertical dispersion coefficients assuming a ratio of 1/10 and 1/100.

The model had a negligible to moderate sensitivity to changes in dispersivity values (**Table 5-3**). The greatest sensitivity for dispersivity was moderate for the highest dispersivity sensitivity test at monitoring well locations G313, G314, G316, and G317. Sensitivity was also moderate for the lower dispersivity sensitivity test at monitoring well locations G313 and G317.

5.2.3.8 Retardation and Decay

It was assumed that sulfate would not significantly sorb or chemically react with aquifer solids (Kd was set to 0 mL/g) which is a conservative estimate for estimating contaminant transport times. Boron, sulfate, and TDS transport is likely to be affected by both chemical and physical attenuation mechanisms (*i.e.,* adsorption and/or precipitation reactions as well as dilution and dispersion). Batch adsorption testing was conducted to generate site specific partition coefficient results for boron and sulfate (Geosyntec, 2022b; **Appendix C**) for locations G311 and G313. Results of the testing are summarized below:

- Boron: The Freundlich isotherm (K_F) fit the data best for G313/SB306 and G313/SB313, yielding K_F values of 0.65 liters per kilogram (L/kg) and 2.03 L/kg, respectively. Though slightly higher at G313/SB313, these values are comparable to boron partition coefficients reported in literature, which range from 0.19 to 1.3 L/kg depending on pH conditions and the amount of sorbent present (EPRI, 2005; Strenge & Peterson, 1989). No partition coefficient was calculated for G311.
- Sulfate: The G311 partition coefficient for sulfate ranged from -624 L/kg for the Langmuir isotherm (K_L) to 10.11 L/kg for the linear isotherm (K_D) , but the best-fitting Freundlich isotherm yielded a low K_F value of 9.2 \times 10⁻¹² L/kg. None of the isotherms showed a high goodness-of-fit (*i.e.,* R2) for either G313/SB306 or G313/SB313, with the highest correlation being 0.05, and were associated with erroneously high (1,700 L/kg) and low (-690 L/kg) partition coefficients. An accurate sulfate partition coefficient could therefore not be calculated from any of the data. These results are consistent with the findings of Strenge and Peterson

(1989), who found that partition coefficients for sulfate are 0.0 L/kg, regardless of pH conditions and the amount of sorbent present.

The results from site samples are variable with poor goodness of fit which supports modeling sulfate without retardation. The potential exceedances identified in groundwater (boron, sulfate, and TDS) are affected by natural attenuation processes in multiple ways and to varying degrees. Further assessment of these processes and how they may be applied as a potential groundwater remedy will be completed as part of future remedy selection evaluations, as necessary. For the purposes of this GMR, and as mentioned at the beginning of this section, no retardation was applied to sulfate transport in the model (*i.e.*, Kd was set to 0 mL/g). Sensitivity tests were not run for retardation.

5.3 Flow and Transport Model Assumptions and Limitations

Simplifying assumptions were made while developing this model:

- Leading up to 2022, the groundwater flow system cannot be simulated as steady state.
- Natural recharge is constant over the long term.
- Fluctuations in lake stage do not affect groundwater flow and transport over the long term.
- Hydraulic conductivity is consistent within hydrostratigraphic units.
- The approximate base of ash surface in the AP1, GMF GSP, GMF RP, and LF were developed with Golder using soil borings and historic topographic maps.
- Source concentrations are assumed to remain constant over time.
- Sulfate is not adsorbed and does not decay and mixing and dispersion are the only attenuation mechanisms.

The model is limited by the data used for calibration, which adequately define the local groundwater flow system and the source and extent of the plume. Since data used for calibration are located near the units on site, model predictions of transport distant spatially and temporally from the calibrated conditions at the CCR units will not be as reliable as predictions closer to the CCR units and concentrations observed in 2021.

5.4 Calibration Flow Model

The groundwater model was manually calibrated to best approximate the mean groundwater elevations in 95 wells at the site. The mean elevations used for calibration and locations of wells within the flow model are summarized in **Table 4-1** Well locations are shown in **Figure 2-1**. This involved modifying the hydraulic conductivities of the different hydrostratigraphic units, recharge rate, and conductance of the drains, rivers, and general head boundaries within the model to minimize the difference between the mean observed groundwater elevation and simulated groundwater elevation. Where possible, the range of the parameter values used during calibration were based on observed values (*i.e.,* for the range in hydraulic conductivity estimates from the HCR). Where this was not possible, such as for the drain and general head boundary conductance, the range of parameter values were based on other site information or inferred from knowledge from similar sites. Where data were limited, the parameter values were less constrained during calibration (*e.g.,* parameter values had wider ranges). The SSR was used as a metric to identify the optimal values for the different parameters. piplied to sulfate transport in the model (*i.e.*, Kd was set to 0 mL/g). Sensitivity tests we

n for retardation.
 3 Flow and Transport Model Assumptions and Limitations
 3 Flow and Transport Model Assumptions and Lim
5.5 Calibration Flow and Transport Model Results

Results of the MODFLOW modeling are presented below. The model files accompany this report (**Appendix B**). **Table 5-1** shows the calibrated hydraulic conductivity for the different units shown in **Figures 5-12 through 5-16**.

Groundwater model calibration results are presented in **Figure 5-22** and **Figure 5-23**, which shows the observed and simulated groundwater elevations and the observed groundwater elevations versus residuals. The near-linear relationship between observed and simulated values presented on **Figure 5-22** indicates that the model adequately represents the calibration dataset. The root mean squared error of the groundwater elevation across all wells was 1.92 feet. The mass balance error for the flow model was 0.00 percent and the ratio of the residual standard deviation to the range of heads was 9.0 percent, which is below the desired target value of 10 percent. Another flow model calibration goal is that residuals are evenly distributed such that there is no bias affecting modeled flow. The observed heads are plotted versus the simulated heads in **Figure 5-23** and simulated values are evenly distributed above and below observed values. The residual mean was also near zero with a value of 0.10 feet, indicating a small bias towards underestimating the groundwater elevations in the calibrated model; this is also illustrated in the observed versus residuals plot in **Figure 5-23**. evations versus residuals. The near-linear relationship between observed and simulated
essented on **Figure 5-22** indicates that the model adequately represents the calibration
tataset. The root mean spare error of the grou

The simulated groundwater elevations within the UA (layer 3) for the entire site are shown in **Figure 5-24**. **Figure 5-25** shows the simulated groundwater elevations in proximity to AP1. In general, the model is able to simulate the groundwater flow patterns for the UA (**Figure 2-2** and **Figure 2-3**) at AP1 as interpreted from the site well data for April and July 2021, respectively. The simulated groundwater flow pattern also captures the radial flow pattern centered on the southwest area of AP1. Fourteen wells provided calibration targets for the simulated groundwater level around AP1. The simulated groundwater levels for five of these wells are within 1 foot; six wells are within 2 feet. G303 and G312 are underestimated by 2.14 feet and 3.06 feet respectively, and G309 is overestimated by 2.24 feet.

The range of observed sulfate concentrations for transport calibration locations are summarized in **Table 4-1**. The goals of the transport model calibration were to have predicted concentrations fall within the range of observed concentrations, and to have predicted concentrations above and below the GWPS for sulfate (400 mg/L) match observed concentrations above or below the standard at each well. One or both of these goals were achieved at all of the transport calibration location wells, except G317, where concentrations were underpredicted (**Figure 5-26**). Deviations from the observed ranges are discussed below.

The model underpredicts concentrations at G305 and G317. The observed sulfate concentrations range from 710 to 930 mg/L and 780 to 1100 mg/L for G305 and G317, respectively. The predicted concentrations are 424.8 mg/L and 146.8 mg/L for G305 and G317, respectively. G305 is located south of AP1 (**Figure 2-1**) in close proximity to the mine entrance discussed in **Section 2.2.4** and shown in **Figure 1-2**. The disturbance associated with the former mining activity may be associated with the elevated sulfate concentrations in this well. G317 is located southeast of AP1, downgradient of G303 (whose predicted sulfate concentration is within the observed range). Groundwater flow in this area is predominantly towards Coffeen Lake (west to east). There is aerial and topographic evidence supporting the presence of a soil pile related to the mining activities in the area west (upgradient) of G317 (see **Section 2.2.4**). One soil boring completed through the soil pile documents the presence of coal in the boring log, indicating the

soil pile may be another source of sulfate. This soil pile may potentially leach sulfate into the groundwater thereby increasing the sulfate concentration at G317 above that which would be attributed to AP1 alone.

The remaining calibration locations had predicted concentrations that fall within the range of observed concentrations and/or have predicted concentrations above and below the GWPS for sulfate (400 mg/L) matching observed concentrations above or below the standard at each well. In other words, there was a very good match between predicted and observed sulfate concentrations relative to wells with concentrations above and below the GWPS. The transport model has achieved a very good calibration using a sulfate source concentration of 1,000 mg/L, even though some wells have observed concentrations that are greater than the source concentration used. The distribution of sulfate concentrations in the calibrated model are presented on **Figure 5-27**. once interactions relative to well, your interactions above and below the GWPS. The transformations relative to wells with concentrations above and below the GWPS. The transformation of such a concentration of such a conce

6. PREDICTIVE SIMULATIONS

6.1 Overview and Prediction Model Development

Prediction simulations were performed to evaluate the effects of closure (source control measures) for AP1 on groundwater quality. The prediction simulations evaluated changes in groundwater sulfate concentrations from Scenario 1: CIP (removal of CCR from the eastern portion of AP1 and consolidation into the western portion of the AP1) and Scenario 2: CBR (removal of all CCR material from AP1). As discussed in **Section 5.2.3.7** physical attenuation (dilution and dispersion) of contaminants in groundwater is simulated in MT3DMS, which captures the physical process of natural attenuation as part of corrective actions for both closure scenarios simulated. No retardation was applied to sulfate transport in the model (*i.e.,* Kd was set to 0 mL/g) as discussed in **Section 5.2.3.8**. ortion of AP1 and consolidation into the western portion of the AP1) and Scenario 2: CBI
emoval of all CCR material from AP1). As discussed in **Section 5.2.3.7** physical attenuation and dispersion) of contaminants in grou

Closure scenarios were simulated by initially removing free liquids from the CCR material over the course of 2 years by placing drain cells within AP1 with an elevation of 618 feet and applying zero recharge to simulate dewatering of the CCR units.

HELP-calculated percolation rates, based on removal and final soil backfill grading designs provided in the Draft CCR Final Closure Plans for Coffeen AP1, GMF GSP, and GMF RP (Golder, 2022), were applied for the different closure scenarios. HELP modeling input and output values are summarized in **Table 6-1** and described in detail below.

The CIP and CBR scenarios were simulated for a 100-year period. The following simplifying assumptions were made during the simulations:

- Removal of free liquids from CCR takes place prior to the CIP and CBR closure scenarios. Drain cells were placed within the units to simulate the removal of free water within the ponds; and recharge was set to zero.
- In the CIP and CBR closure scenarios, HELP-calculated average annual percolation rates were developed from a 30-year HELP model run. This 30-year HELP-calculated percolation rate remained constant over duration of the closure scenario prediction model runs following CCR dewatering period.
- Changes in recharge resulting from removal of free liquids (decrease calibration model recharge rates to zero) and CCR fill removal/final soil backfill grading (recharge rates are based on HELP-calculated average annual percolation rates) have an instantaneous effect on recharge and percolation through surface materials.
- Sulfate source concentrations were assumed to be negligible (0 mg/L) in CCR removal areas in both the CIP and CBR scenarios. The spatial distribution of CCR concentrations within the consolidation area for the CIP scenario were maintained from the initial transport simulation.
- Cap construction in CIP scenario was assumed to be completed with a cover system consisting of the following (listed from ground surface down): a vegetative cover (6 inches thick), rooting zone (18 inches thick), a 200-mil geocomposite drainage layer and a 40-mil linear HDPE geomembrane.
- The start of each closure prediction simulation was initiated at the end of the calibration model period of 42 years plus 2 years to complete removal of free liquids. For example, the

simulation of Scenario 1: CIP begins at 44 years (42 years for calibration plus 2 years). The prediction modeling timeline for each scenario is illustrated in **Figure 4-1**.

- CCR consolidation/removal areas were assumed to be graded and include proper drainage controls to remove excess water from the surface using the design drawings provided (Golder, 2022).
- The CIP scenario includes the placement of a stormwater pond within the removal area. The outflow elevation of this stormwater pond is 625 feet, which will discharge into Coffeen Lake adjacent to the AP2. This is represented as a drain in the model whose elevation is equal to the stormwater pond outflow elevation.
- Local fill materials applied to the prediction models have similar hydraulic properties as the UCU materials used in the transport calibration models. However, the local fill materials were assumed to have reduced vertical anisotropy ratios, approaching isotropic, due to reworking of the material as it is placed as backfill (Kh/Kv decreased from measured values of 10 to 1 for reworked material).

6.2 HELP Model Setup and Results

HELP (Version 4.0; Tolaymat and Krause, 2020) was used to estimate percolation through AP1 in areas of CCR removal with soil backfill, and areas of CCR consolidation with final cover system. HELP input and output files are included electronically and attached to this report.

HELP input data and results are provided in **Table 6-1**. All scenarios were modeled for a period of 30 years. Climatic inputs were synthetically generated using default equations developed for Belleville Scott Air Force Base in Belleville, Illinois (the closest weather station included in the HELP database). Precipitation, temperature, and solar radiation was simulated based on the latitude of CPP. Thickness of soil backfill and soil runoff input parameters were developed for the ash fill removal scenarios using data provided in the Draft CCR Final Closure Plans for Coffeen AP1, GMF GSP, and GMF RP (Golder, 2022).

HELP model results (**Table 6-1**) indicated 7.85 inches of percolation per year for AP1 CCR removal and soil backfill area in the CIP scenario and 0.00027 inches of percolation per year through the CCR and final cover system for the CIP scenario. Results indicated 7.85 inches and 6.28 inches of percolation per year for AP1 eastern and western CCR removal and soil backfill area in the CBR scenario, respectively. The differences in HELP model runs for each area included the following parameters: area, soil backfill thickness, slopes, and soil runoff slope length; all other HELP model input parameters were the same for each simulated area. HELP input data and results are provided in **Appendix B.** outflow elevation of this stormwater pond is 625 feet, which will discharge into Coffeen
adjacent to the AP2. This is represented as a drain in the model whose elevation is equ
the stormwater pond outflow elevation.
Local

6.3 Simulation of Closure Scenarios

The calibrated model was used to evaluate the effectiveness of the two closure scenarios by defining CCR removal and consolidation areas, reducing head to simulate removal of free liquids, removing source concentrations from the removal areas, adding drain cells and removing recharge to simulate stormwater management within the removal areas, and applying reduced recharge in the CCR consolidation areas to simulate the effects of the cover system on flow and transport. Removal of source inputs from the ash removal areas was simulated by reducing the sulfate concentrations associated with recharge in the areas to 0 mg/L and removing constant concentration cells.

Each prediction scenario was simulated as a continuation of the AP1 dewatering simulation which followed the transient calibrated model. The prediction model input values are summarized in **Table 6-2,** and the modifications to the recharge zones and drain placement for the CIP scenario are illustrated in **Figure 6-1**. **Figure 6-2** illustrates the CCR removal area for the CBR at the AP1. The two closure scenarios are discussed in this report based on predicted changes in sulfate concentrations as described below and results are presented in **Figure 6-3 to Figure 6-6**.

6.3.1 Closure in Place Model Results

The design for Scenario 1: CIP includes an initial 2-year dewatering period to remove free liquids followed by CCR removal from AP1, consolidation in the western area of AP1, and construction of a cover system over the remaining CCR (**Figure 6-1**). Stormwater drainage will be present within the eastern area of AP1 with an outflow elevation of 625 feet.

Predicted concentrations start to decline at all monitoring wells with observations above the GWPS for sulfate (400 mg/L) once closure actions are initiated within the prediction model. These declines occur first in the eastern area where CCR is removed and saturated ash cells (constant concentration cells) are reduced in the area of the highest modeled source concentrations. Following removal of CCR in the eastern area, sulfate concentrations are no longer entering the model domain from recharge or from saturated ash cells (constant concentration cells). Dewatering also reduces the head within AP1. These low heads are maintained following completion of closure by the drain cells that simulate storm water management designs within the removal area to the east, and by the greatly reduced infiltration rates (recharge) that result from placement of the cover system over the consolidated CCR in the western end of AP1. As a result of the reduced heads and recharge, downward percolation of solute mass from AP1 is reduced, which decreases the sulfate concentration entering the model domain. **S.1 CIOSURE IN PIAGE MOOEI NESSIDE**

and exaign for Scenario 1: CIP includes an initial 2-year dewatering period to remove free design for Scenario 1: CIP includes an initial 2-year dewatering period to remove free fourc

The predictive model indicates that most wells will reach the GWPS (400 mg/L) in under 14.8 years following closure, with one exception. **Figure 6-3** and **Figure 6-4** show the extent of the plume in the UA after 14.8 years and the maximum extent of the plume in the model after 14.8 years, respectively. The predicted delayed reduction in concentration at well G301 is a result of the well's location along the flow path of the residual sulfate concentrations released into native geologic materials prior to closure. All UA groundwater monitoring wells are below the GWPS within 58.8 years (**Figure 6-5** and **Figure 6-6**). The residual sulfate plume in the UA from the calibrated model remains in close proximity to AP1 as it recedes over time. The predicted footprint of the sulfate plume in the UA after 58.8 years shown in **Figure 6-5** is considerably reduced from that at the end of the transient model simulation (**Figure 5-27**).

The predicted delayed reduction in concentration at well G301 is a result of the well's location along the flow path of the residual sulfate concentrations released into native geologic materials prior to closure. Reduced percolation rates through the consolidation area within AP1 in the CIP scenario means that the residual sulfate concentrations require a longer time period to migrate through native geologic materials.

6.3.2 Closure by Removal Model Results

The design for Scenario 2: CBR includes an initial 2-year dewatering period followed by CCR removal from AP1 (**Figure 6-2**). Stormwater drainage is present within AP1 with an outflow elevation of 625 feet.

For most wells, predicted concentrations for CBR start to decline at monitoring wells with observations above the standard GWPS for sulfate (400 mg/L) once the closure actions are initiated within the prediction model. The concentration of sulfate in some wells (most notably G315, G307 and G308) show short term fluctuations (less than 5 years) following the removal of concentration during the dewatering phase, such that sulfate concentrations decline and are followed by a short rise before the impacts of the CBR are clearly observed. The general decline in sulfate concentration occur as the CCR is removed from AP1 and saturated ash cells (constant concentration cells) are removed. Following removal of CCR, sulfate concentrations are no longer entering the model domain from recharge or from saturated ash cells (constant concentration cells); all source concentrations are removed. Dewatering through removal of free liquids also reduces the head within AP1. These low heads are maintained following completion of closure by the drain cells that simulate storm water management designs within AP1. The removal of the CCR sources leads to the gradual reduction the residual sulfate concentrations released into native geologic materials prior to closure. All monitoring wells with observations above the standard GWPS for sulfate (400 mg/L) are predicted to be below the GWPS 15.4 years after closure implementation (**Figure 6-3**).

The sulfate plume in the CBR prediction model differs from that in the CIP prediction model. Higher recharge rates are present in the western portion of the pond because there is no cover system. The relatively higher recharge rates maintain components of the radial flow pattern described in **Section 2.2** at AP1. However, the stormwater drainage within the pond does constrain the groundwater elevation beneath AP1. As a result of the radial flow pattern, the prediction model indicates that a portion of the historic plume will remain along the western edge of AP1 as the plume recedes over time. The maximum extent of the plume at 14.8 years is illustrated in **Figure 6-4**. The maximum extent of the plume remains in close proximity to AP1 and is no longer present above the GWPS (400 mg/L) at 58.8 years as illustrated in **Figure 6-6**. meteriorion cells) are removed. Following removal of CCR, sulfate concentrations are nucleated in the medel domain from recharge or from saturated ash cells (constant concentrations are removed. Devatering through removal

7. CONCLUSIONS

This GMR has been prepared to evaluate how proposed CIP and CBR scenarios will achieve compliance with the applicable groundwater standards at the CPP for AP1. An existing groundwater model was updated to include data collected from the recent 2021 field investigations and used to predict the impacts of the closure scenarios on groundwater quality at the CPP. Statistically significant correlations between sulfate concentrations and concentrations of TDS identified as potential exceedances of the GWPS indicate sulfate is an acceptable surrogate for TDS in the groundwater model. Concentrations of TDS are expected to change along with model predicted sulfate concentrations. A potential exceedance of boron was observed at one monitoring well, G313, which also has potential exceedances of both sulfate and TDS. Similar source and behavior in the groundwater system would be expected among boron, sulfate, and TDS at UA monitoring well G313, and boron concentrations are expected to change along with model predicted sulfate concentrations. It was assumed that sulfate would not significantly sorb or chemically react with aquifer solids (Kd was set to 0 mL/g) which is a conservative estimate for predicting contaminant transport times in the model. The MODFLOW and MT3DMS models were used to evaluate two scenarios using information provided in the Draft CCR Final Closure Plan (Golder, 2022): Control and the groundwater model. Concentrations of the GWPS indicate sulfate is an acceptable sur-
To TDS in the groundwater model. Concentrations of TDS are expected to change along
y deterified as potential exceedances

- **Scenario 1:** CIP including removal of CCR from the eastern portion of AP1, consolidation into the western portion of AP1, and construction of a cover system over the remaining CCR.
- **Scenario 2:** CBR including removal of all CCR and regrading of the removal area.

Differences exist in the timeframes to reach the GWPS for most monitoring wells between CIP and CBR. In general, the simulated groundwater concentrations in the monitoring wells within the UA will achieve the GWPS in 15 years for both the CIP and CBR closure scenarios, with the exception of well G301 in the CIP scenario. The predicted delayed reduction in concentration at well G301 is a result of the well's location along the flow path of the residual sulfate concentrations released into native geologic materials prior to closure. Reduced percolation rates through the consolidation area within AP1 in the CIP scenario means that the residual sulfate concentrations require a longer time period to migrate through native geologic materials.

Results of groundwater fate and transport modeling conservatively estimate that groundwater concentrations will attain the GWPS for all constituents identified as potential exceedances of the GWPS in the UA monitoring wells within 59 years of closure implementation for CIP and 15 years for CBR. The residual sulfate plumes from the calibrated model remain in close proximity to AP1 and has been simulated to decline below the GWPS (400 mg/L) within 59 years for CBR. The residual plume in the CIP scenario will take longer in a small area at the northwest corner of AP1 due to the reduced infiltration rates below the cover system.

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GROUNDWATER MODELING REPORT COFFEEN POWER PLANT ASH POND NO. 1 COFFEEN, ILLINOIS

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GROUNDWATER MODELING REPORT COFFEEN POWER PLANT ASH POND NO. 1 COFFEEN, ILLINOIS

Notes:

All elevation data are presented relative to the North American Vertical Datum 1988 (NAVD88), GEOID 12A

-- = data not available

bgs = below ground surface

DA = deep aquifer

ft = foot or feet

HSU = hydrostratigraphic Unit

LCU = lower confining unit

PVC = polyvinyl chloride

S = source water

SW = surface water

UA = uppermost aquifer

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TABLE 4-1. FLOW AND TRANSPORT MODEL CALIBRATION TARGETS

GROUNDWATER MODELING REPORT

COFFEEN POWER PLANT

ASH POND NO. 1

COFFEEN, ILLINOIS

TABLE 4-1. FLOW AND TRANSPORT MODEL CALIBRATION TARGETS

GROUNDWATER MODELING REPORT

COFFEEN POWER PLANT

ASH POND NO. 1

COFFEEN, ILLINOIS

TABLE 4-1. FLOW AND TRANSPORT MODEL CALIBRATION TARGETS

GROUNDWATER MODELING REPORT

COFFEEN POWER PLANT

ASH POND NO. 1

COFFEEN, ILLINOIS

AP1 = Ash Pond No. 1 CCR = coal combustion residuals AP2 = Ash Pond No. 2 and 2 uppermost aquifer UA = uppermost aquifer

B = Background UCU = lower confining un $LCU = lower$ confining unit

Notes:

GSP = Gypsum Management Facility Gypsum Stack Pond

LF = Landfill

max=maximum

mg/l = milligrams per liter

min=minimum

RP = Gypsum Management Facility Recycle Pond

std=standard deviation from the mean

HSU = Hydrostratigraphic Unit

GROUNDWATER MODELING REPORTCOFFEEN POWER PLANT ASH POND NO. 1 COFFEEN, ILLINOIS

Notes:

¹ Sensitivity Explanation:

Negligible - SSR changed by less than 1% **And the set of the set of the set of the Hydrostratigraphic Unit**
Low - SSR change between 1% and 10% **And the set of the**

Low - SSR change between 1% and 10% and 50% UCU = upper confining unit UCU = upper conf

Moderate - SSR change between 10% and 50% UA = uppermost aquifer UA = uppermost aquifer
Moderately High - SSR change between 50% and 100% UA = UACU = lower confining unit Moderately High - SSR change between 50% and 100%

GSP = Gypsum Management Facility Gypsum Stack Pond $in/yr = inches per year$

High - SSR change greater than 100%

² Liner thickness accounted for in harmonic mean calculation
SSR = sum of squared residuals

 $- - = \text{not tested}$

 $AP1 =$ Ash Pond No. 1

AP2 = Ash Pond No. 2

CCR = coal combustion residualscm/s = centimeters per second

 $ft/d = feet per day$

ft²/day = feet squared per day

Kh/Kv = anisotropy ratio

LF = Landfill

NA = not applicable

RP = Gypsum Management Facility Recycle Pond

SW = Surface Water

TABLE 5-2. TRANSPORT MODEL INPUT VALUES (CALIBRATION)

GROUNDWATER MODELING REPORT COFFEEN POWER PLANT ASH POND NO. 1 COFFEEN, ILLINOIS

TABLE 5-2. TRANSPORT MODEL INPUT VALUES (CALIBRATION)

GROUNDWATER MODELING REPORT COFFEEN POWER PLANT ASH POND NO. 1 COFFEEN, ILLINOIS

[O: SLN 04/01/22; C: EGP 04/29/22]

Notes:

 $- - =$ not tested

 $AP1 =$ Ash Pond No. 1

Hydrostratigraphic Unit

UCU = upper confining unit

AP2 = Ash Pond No. 2

 UA = uppermost aquifer LCU = lower confining unit

CCR = coal combustion residuals

ft/d = feet per day

GMF = Gypsum Management Facility

GSP = Gypsum Management Facility Gypsum Stack Pond

LF = Landfill

mg/L = milligrams per liter

NA = not applicable

RP = Gypsum Management Facility Recycle Pond

 $SS = Steady State model$

STP = Stress Period

SW = Surface Water

TR = Transient model

TABLE 5-3. TRANSPORT MODEL INPUT SENSITIVITY (CALIBRATION)

GROUNDWATER MODELING REPORT COFFEEN POWER PLANT

ASH POND NO. 1

COFFEEN, ILLINOIS

TABLE 5-3. TRANSPORT MODEL INPUT SENSITIVITY (CALIBRATION)

GROUNDWATER MODELING REPORT

COFFEEN POWER PLANT

ASH POND NO. 1

COFFEEN, ILLINOIS

Notes:

¹ Sensitivity Explanation:

Negligible = concentration changed by less than 1%

[O: SLN 04/10/22; C: EGP 5/5/22]

Low = concentration change between 1% and 10%

Moderate = concentration change between 10% and 50%

Moderately High = concentration change between 50% and 100%

High = concentration change greater than 100%

² sensitivity test used transient transport

AP1 = Ash Pond No. 1

AP2 = Ash Pond No. 2

Disp = dispersivity

GSP = Gypsum Management Facility Gypsum Stack Pond

HFB = Horizontal Flow Boundary

ID = identification

mg/L = milligrams per liter

RP = Gypsum Management Facility Recycle Pond

 $S =$ storativity

Sy = specific yield

TABLE 6-1. HELP MODEL INPUT AND OUTPUT VALUES

TABLE 6-1. HELP MODEL INPUT AND OUTPUT VALUES

TABLE 6-1. HELP MODEL INPUT AND OUTPUT VALUES

GROUNDWATER MODELING REPORT COFFEEN POWER PLANTASH POND NO. 1 COFFEEN, ILLINOIS

Notes:

 $% =$ percent

 $CBR = closure by removal$

 $CIP = closure$ in place

cm/s = centimeters per second

 $ft = feet$

HCR = Hydrogeologic Site Characterization Report

HELP = Hydrologic Evaluation of Landfill Performance

in = inches

in/yr = inches per year

Lat = latitude

Long = longitude

References:

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TABLE 6-2. PREDICTION MODEL INPUT VALUES

GROUNDWATER MODELING REPORT COFFEEN POWER PLANT ASH POND NO. 1COFFEEN, ILLINOIS

Notes:

 $- - -$ = not included

 $AP1 =$ Ash Pond No. 1

CCR = coal combustion residuals

ft/day = feet per day

inches/yr = inches per year

mg/L = milligrams per liter

[O: SLN 04/01/22; C: EGP 04/29/22]

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FIGURE 4-1

Rocky Ford Lakes Rocky Ford Sportsman Club North Lake Coffeen Lake LANDFILL GMF GYPSUM STACK POND GMF RECYCLE POND ASH POND 2 ASH POND 1 GSP GRP CONNECTOR CHANNEL Unnamed Tributary West Lake Coffeen Unnamed Tributary East Lake Coffeen SURFACE WATER POND COOLING POND PROJECT: 169000XXXX | DATED: 5/6/2022 | DESIGNER: gal (N $\overline{\mathsf{N}}$ Y:\Mapping\Projects\22\2285\MXD\Model_Figures\Coffeen\AP1\Figure 5-1_Model Area Map. *COFFEEN LAKE*

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RAMBOLL

MODEL RECHARGE DISTRIBUTION STEADY STATE (SS) MODEL

GROUNDWATER MODELING REPORT ASH POND NO. 1 COFFEEN POWER PLANT COFFEEN, ILLINOIS MODEL RECHARGE DISTRIBUTION STEADY STATE (SS) MODEL

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SIMULATED GROUNDWATER LEVEL RESIDUALS FROM THE CALIBRATED MODEL

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FROM UA (LAYER 3) FROM THE CALIBRATED MODEL

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RAMBCLL

SIMULATED SULFATE PLUME IN THE UA FROM THE TRANSIENT MODEL

GROUNDWATER MODELING REPORT ASH POND NO. 1 COFFEEN POWER PLANT COFFEEN, ILLINOIS SIMULATED SULFATE PLUME IN THE UA FROM THE TRANSIENT MORE CROUNDWATER MODELING REPORT

GROUNDWATER MODELING REPORT

ASH POND NO. 1

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COFFEEN, ILLINOIS

SIMULATED SULFATE PLUME OF THE UA FOR THE CIP AND CBR SCENARIOS AFTER 14.8

YEARS

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SIMULATED MAXIMUM EXTENT OF THE SULFATE PLUME FOR THE CIP AND CBR

SCENARIOS AFTER 14.8 YEARS

GROUNDWATER MODELING REPORT

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COFFEEN, ILLINOIS SCENARIOS AFTER 14.8 YEARS

SIMULATED SULFATE PLUME OF THE UA FOR THE CIP AND CBR SCENARIOS AFTER 58.8

YEARS

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SIMULATED MAXIMUM EXTENT OF THE SULFATE PLUME FOR THE CIP AND CBR

SCENARIOS AFTER 58.8 YEARS

GROUNDWATER MODELING REPORT

ASH POND NO. 1

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COFFEEN, ILLINOIS SCENARIOS AFTER 58.8 YEARS

APPENDICES

RAFT

APPENDIX A EVALUATION OF POTENTIAL GWPS EXCEEDANCES (GEOSYNTEC CONSULTANTS, INC., 2022) ENDIX A
LUATION OF POTENTIAL GWPS EXCEEDANCES
SYNTEC CONSULTANTS, INC., 2022)

DRAFT EVALUATION OF POTENTIAL GROUNDWATER PROTECTION STANDARD EXCEEDANCES

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IMcBride and Son Center Dr Su

Chesterfield, Missouri

Prepared for

Illinois Power Generating Company

Submitted by

1 McBride and Son Center Dr Suite 202 Chesterfield, Missouri 63005

May 6, 2022

GLP8029
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LIST OF ACRONYMS AND ABBREVIATIONS

SECTION 1

INTRODUCTION

The Illinois Power Generation Company (IPGC) currently operates the Coffeen Power Plant (CPP) and its associated ash ponds. In October 2021, the IPGC submitted an Operating Permit application for the coal combustion residual (CCR) Unit referred to as the Ash Pond Number (No.) 1 (AP1), Vistra identification (ID) No. 101, IEPA ID No. W1350150004-01, and National Inventory of Dams (NID) No. IL50722 (Burns & McDonnell, 2021). The Operating Permit was prepared to comply with Part 845 "Standards of the Disposal of Coal Combustion Residuals in Surface Impoundments (Part 845), which was promulgated by the Illinois Pollution Control Board on April 21, 2021. Ramboll Americas Engineer Solutions, Inc. (Ramboll) identified potential groundwater protection standard (GWPS) exceedances for multiple constituents in groundwater samples collected from monitoring wells in the vicinity of AP1, as presented in the Operating Permit Application. This report was developed to further evaluate the potential GWPS exceedances identified. coal combustion residual (CCR) Unit referred to as the Ash Pond Number (No.) I
dentification (ID) No. 101, IEPA ID No. W1350150004-01, and National Inven
NID) No. ILS0722 (Burns & McDonnell, 2021). The Operating Permit was

SECTION 2

BACKGROUND

A brief description of the site location, AP1 design, geology, and groundwater assessment activities to date are described below.

2.1 Site Location and Description

The CPP, operated by the IPGC is located in Montgomery County, Illinois approximately two miles south of the City of Coffeen in Section 11, Township 7 North, and Range 7 East. The CPP is located between the two lobes of Coffeen Lake to the west, east, and south, and is bordered by agricultural land to the north. The CPP operated as a coal-fired power plant from 1964 to November 2019 and has five CCR management units. The approximately 1,100-acre Coffeen Lake was built by damming the McDavid Branch of the East Fork of Shoal Creek in 1963 for use as an artificial cooling lake for the CPP. Historically, coal mines were operated at depth in the vicinity of the CPP as well as a US Minerals processing facility located to the north. Mine shafts, processing facilities, and coal storage associated with these historical operations were located south of AP1. **Example 12**
 Example 12 Example 12

2.2 Ash Pond 1 Design

Coffeen AP1 is a 23-acre, unlined surface impoundment used to manage CCR and non-CCR waste streams at the CPP. The location of AP1 relative to the proposed monitoring well network is displayed on Figure 2-1 of the Groundwater Monitoring Plan Report (Ramboll, 2021a) and is provided herein as **Appendix A**. AP1 (also known as the Bottom Ash/Recycle Pond) is a reclaimed ash pond that was constructed utilizing the existing earthen berms with reinforcement. AP1 is an unlined surface impoundment which covers an area of approximately 23 acres, has berms up to 41 feet above the surrounding land surface, and a capacity of 300 acre-feet. It primarily received bottom ash and low volume wastes from floor drains in the main power block building. Several years ago, air heater wash and boiler chemical cleaning wastes were directed to AP1, but this practice was discontinued. The bottom ash is periodically removed from AP1 for beneficial uses by a third-party contractor. Sluicing of waste to AP1 ceased prior to November 4, 2019.

2.3 Geology and Hydrogeology

The AP1 geologic and hydrogeologic setting summarized below is excerpted from the Hydrogeologic Site Characterization Report (HCR) for AP1 (Ramboll, 2021b).

There are five principal layers of unlithified material present below AP1 and above bedrock which are categorized into hydrostratigraphic units listed below (from the surface downward) based on stratigraphic relationships and hydrogeologic characteristics:

- **Upper Confining Unit (UCU):** Composed of the Roxana and Peoria Silts (Loess Unit) and the upper clayey portion of the Hagarstown member which are classified as silts-clayey silts and gravelly clay below the surficial soil. The UCU has been eroded east of AP1, near the Unnamed Tributary.
- **Uppermost Aquifer (UA):** The uppermost aquifer is the Hagarstown Member which is classified as primarily sandy-gravelly silts and clays with thin beds of sands. Similar to the Loess Unit, the Hagarstown is absent in some locations near the Unnamed Tributary.
- **Lower Confining Unit (LCU):** Comprised of the Vandalia Member, Mulberry Grove Member, and Smithboro Member. These units include a sandy-silty till with thin, discontinuous sand lenses, a discontinuous and limited extent sandy silt which has infilled prior erosional features, and silty-clayey diamicton, respectively. This unit has been identified as a potential migration pathway (PMP) because downward vertical gradients indicate that there is the potential for impacts to migrate within this unit. gravelly elay below the surficial soil. The UCU has been croded east of AP1, n
anned Tributary.

Dermost Aquifer (UA): The uppermost aquifer is the Hagarstown Member w

sified as primarily sandy-gravelly silts and clays wi
- **Deep Aquifer (DA):** Comprised of sand and sandy silt/clay units of the Yarmouth Soil, which include accretionary deposits of fine sediment and organic materials, typically less than five feet thick and discontinuous across the CPP. This unit is also identified as a PMP, because it is the first permeable unit below the uppermost aquifer.
- **Deep Confining Unit (DCU):** Comprised of the Banner Formation, generally consisting of clays, silts, and sands. The Lierle Clay Member is the upper layer of the Banner Formation which was encountered at the Site.

Bedrock is comprised of the Bond Formation, which consists of limestone and calcareous clays and shale. Bedrock was not encountered in the borings advanced to date at CPP.

Flow of groundwater from central portions of the CPP to Coffeen Lake or the Unnamed Tributary through the UA are the primary pathways for contaminant migration. The LCU and DA underlying the UA have been identified as PMPs. Groundwater elevations are primarily controlled by surface topography, geologic unit topography, and water levels within Coffeen Lake and the Unnamed Tributary. A groundwater divide trending north-south is observed running through the approximate center of the CPP (Figure 1-3 of Ramboll [2021a], provided as **Appendix B**). Phreatic surfaces or water elevations within the surface impoundments are generally consistent and have

not been observed to fluctuate with groundwater elevations, indicating limited hydraulic connection with the surface impoundments.

2.4 Groundwater and AP1 Monitoring

The proposed Part 845 monitoring well network for AP1 was established in the Groundwater Monitoring Plan (Ramboll, 2021a). The proposed monitoring well network consists of sixteen (16) monitoring wells, which are installed in the UA, LCU, DA, and temporary water-level only surface water staff gages. Two of the installed wells are background monitoring wells (G281 and G306) and the remaining fourteen are compliance monitoring wells. Both background wells and most compliance wells are screened within the UA. G307D, G314, and G316 are screened within the LCU, and G314D is screened within the DA. Well locations are shown on **Appendix A**. ma Pian (Kambol), 2021a). In proposed monitoring well network consists of sixted
ting well, which are installed in the UA, LCU, DA, and temporary vater-level only
farf gages. Two of the installed wells are background monit

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SECTION 3

POTENTIAL GROUNDWATER PROTECTION STANDARD EXCEEDANCE REVIEW

An evaluation of the history of potential GWPS exceedances was completed for the Operating Permit application in October 2021 (Burns & McDonnell, 2021). Groundwater concentrations from 2015 to 2021 were evaluated for potential exceedances in accordance with the Statistical Analysis Plan proposed in the Operating Permit application. Potential exceedances are summarized below:

- Boron at monitoring well G313: The boron statistical result at G313 is 3.5 milligrams per liter (mg/L), which exceeds the Part 845 GWPS (3.2 mg/L).
- Cobalt at monitoring well G314: The cobalt statistical result at G314 is 0.00959 mg/L which exceeds the Part 845 GWPS (0.006 mg/L).
- pH (field) at monitoring well G312: The pH statistical result at G312 is 6.4 standard units (SU), which is below the lower limit of the Part 845 GWPS (6.5/9.0 SU).
- Sulfate at monitoring wells G301, G303, G304, G305, G307, G307D, G308, G309, G310, G311, G312, G313, G314, G314D, G315, and G317: The sulfate statistical results ranged from 464 to 1100 mg/L and individually exceed their relevant Part 845 GWPS (400 to 700 mg/L) for the identified wells. is Plan proposed in the Operating Permit application. Potential execedance
ized below:
Boron at monitoring well G313: The boron statistical result at G313 is 3.5 milligra
liter (mg/L), which exceeds the Part 845 GWPS (3.2
- Total dissolved solids (TDS) at monitoring wells G303, G304, G305, G307, G307D, G308, G309, G310, G311, G312, G313, G314, G315, and G317: The TDS statistical results ranged from 1210 to 1900 mg/L which exceed the Part 845 GWPS (1200 mg/L).

A review of groundwater, porewater, soil, and ash data indicates that the potential exceedances of cobalt at G314 and pH at G312 are not related to AP1, as documented in Section 4. An evaluation of alternative sources of the boron, sulfate, and TDS potential exceedances was not completed at this time.

SECTION 4 LINES OF EVIDENCE

A review of groundwater, porewater, soil, and ash data indicates that the potential GWPS exceedances of cobalt at G314 and the pH value at G312 are not related to AP1, as supported by the lines of evidence (LOE) below:

- 1. AP1 porewater samples do not contain detectable concentrations of cobalt.
- 2. Cobalt concentrations in ash samples collected from AP1 are comparable to or lower than cobalt concentrations in soil samples near AP1.
- 3. Monitoring well G314 has experienced significant changes in oxidation-reduction (redox) conditions since well installation occurred, which may impact cobalt behavior in groundwater.
- 4. AP1 porewater is slightly basic and would not result in low pH measurements at monitoring well G312.
- 5. pH values within the proposed monitoring well network are strongly correlated with saturation indices of carbonate minerals in soil near AP1.

4.1 LOE #1: AP1 porewater samples do not contain detectable concentrations of cobalt

Of the three AP1 porewater sampling locations analyzed for cobalt (AP1d, XPW01, and XPW02), none have ever contained cobalt concentrations above the method detection limit of 0.002 mg/L; therefore, cobalt concentrations detected at G314 cannot be derived from a mixing scenario between groundwater and AP1 porewater. In contrast, both background monitoring wells have at times contained cobalt concentrations within the range observed at G314. This indicates that aqueous cobalt is naturally present in groundwater at CPP at variable concentrations. AP1 porewater samples do not contain detectable concentrations of cobalt.
Cobalt concentrations in ash samples collected from AP1 are comparable to or low
cobalt concentrations in soil samples near AP1.
Monitoring well G31

Figure 1 displays cobalt concentrations over time for G314, background wells G306 and G281, and porewater samples from AP1. Cobalt concentrations at G314 display an increasing trend, but this trend is punctuated by a concentration decrease in the most recent sampling event. The highest values at G314 are comparable to or lower than select results observed at background well G306, suggesting there is variability within the aquifer.

4.2 LOE #2: Cobalt concentrations in ash samples collected from AP1 are comparable to or lower than cobalt concentrations in soil samples near AP1

Soil samples were collected in May 2021 and September 2021 adjacent to select existing monitoring wells and analyzed for total metals. Cobalt concentrations in soil are displayed in **Table 1** along with total cobalt concentrations in ash material collected from AP1. Cobalt concentrations in ash from AP1 $(4.3 - 4.8 \text{ mg/kg})$ fall within the range of cobalt concentrations observed in CPP soil (4.0 – 10 mg/kg). Cobalt concentrations in soil are highest at Ash Pond No.2 (AP2) background monitoring well G270, which is in a background location relative to AP1 (**Appendix B**). **Table 1** indicates variability in cobalt concentrations detected in soil across the CPP. Three sample locations (two background locations and one compliance location) contained greater cobalt concentrations than ash samples, indicating that naturally occurring cobalt exists in solid phase across the CPP at equivalent or greater concentrations than within AP1 itself.

4.3 LOE #3: Monitoring well G314 has experienced significant changes in oxidationreduction (redox) conditions since well installation occurred, which may impact cobalt behavior in groundwater

Groundwater oxidation-reduction potential (ORP) was measured as a field parameter during the sample collection process at monitoring wells in the proposed network. ORP is a measure of the redox conditions of water which, along with other parameters like pH, temperature, and chemical composition, govern the stability of minerals comprising groundwater aquifer solids. ORP values over time at recently installed compliance monitoring wells are displayed on **Figure 2**. ORP values for recently installed wells display a decreasing trend, indicating a shift from highly oxic to near reducing conditions. This decreasing trend is hypothesized to be attributable to stabilization of the new wells following the potential introduction of drilling water involved in the well installation process. Such a change in geochemical conditions can influence the stability of redox-sensitive mineral phases such as iron and manganese oxides. Significantly, decreases in ORP are commonly correlated with dissolution of iron and manganese bearing minerals, leading to the release of ions associated with these mineral phases. from AP1 (4.3 - 4.8 mg/kg) fall within the range of cobalt concentrations observed
0 - 10 mg/kg). Cobalt concentrations in soil are highest at Ash Pond No.2
0 und monitoring well G270, which is in a background location re

Cobalt is known to undergo isomorphic substitution for iron in crystalline iron minerals such as iron oxides, iron sulfides, and iron carbonates due to the similar ionic radii of approximately 1.56 angstroms (Å) for iron vs. 1.52 Å for cobalt (Clementi and Raimondi, 1963; Krupka and Serne, 2002; Hitzman et al., 2017). Soil samples around AP1 were collected and submitted for mineralogical analysis via X-ray diffraction (XRD) to determine the mineralogical composition of the natural aquifer material. XRD results are shown in **Table 2**. **Table 2** indicates that the majority component of site soils consists of geochemically inert minerals quartz and feldspar (microcline and albite). No iron oxides or iron sulfides were detected in XRD analysis, but iron-bearing carbonate mineral ankerite $(Ca, Fe(CO₃)₂)$ was detected at a maximum abundance of 7.7 wt.%.

An Eh-pH diagram displaying the thermodynamic stability of iron phases was generated using the average composition of G314 groundwater (**Figure 3**). Geochemical conditions during initial sampling events favored thermodynamic stability of the ferric (Fe^{3+}) iron hydroxide mineral Fe(OH)3; however, no iron hydroxide or oxide minerals were present in XRD results above the detection limit of 0.5%. **Figure 3** indicates G314 groundwater conditions have shifted in recent sampling events, favoring the formation of ferrous (Fe^{2+}) carbonate mineral siderite ($FeCO₃$). The modeled shift of thermodynamic stability away from iron hydroxide and oxide minerals and towards iron carbonates would result in the release of iron and isomorphically substituted cobalt into groundwater through mineral dissolution reactions.

While siderite was not detected in the XRD results, iron-bearing carbonate mineral ankerite was detected at abundances of up to 7.7 wt.%. Ankerite exists in nature as a solid-solution mineral without a fixed mineral formula. As a result, accurate thermodynamic information is not available for modeling purposes and ankerite was consequently not included in the thermodynamic database used to generate **Figure 3**. It is likely that ankerite thermodynamic stability is favored over siderite stability at G314 and the ankerite detected in XRD analyses is a product of the formation of carbonate minerals in an iron-rich environment.

Naturally occurring cobalt is known to substitute for iron in iron-bearing minerals. Thermodynamic modeling indicates that a recent trend in redox conditions has resulted in a mineral stability shift from iron hydroxides and oxides towards iron carbonates. The presence of ankerite, an iron-bearing carbonate mineral, has been confirmed across the site. The modeled dissolution of iron hydroxide and oxide minerals may have resulted in isomorphically substituted cobalt being released from the crystal structure of these minerals and entering groundwater. The presence of observed iron carbonate minerals in soil samples supports the occurrence of this mineralogical shift. d shift of thermodynamic stability away from iron hydroxide and oxide minera iron carbonates would result in the release of iron and isomorphically substituted undwater through mineral dissolution reactions.

and a fabured

4.4 LOE #4: AP1 porewater is slightly basic and would not result in low pH measurements at monitoring well G312

Groundwater pH conditions were measured as a field parameter during the sample collection process at monitoring wells within the proposed monitoring well network. A time series plot of field pH measurements at G312, background wells G281 and G306, and AP1 porewater monitoring locations XPW-01 and XPW-02 is provided as **Figure 4**. Groundwater at monitoring well G312 contains pH levels below the calculated lower GWPS for pH of 6.5 SU. Low pH values at G312 cannot be attributed to AP1, because AP1 porewater samples are consistently slightly basic (pH values range from 7.78-8.08). Physical mixing of AP1 porewater with G312 groundwater would result in an increase in pH at G312. In contrast, pH values at background well G306 were occasionally measured at 6.5 SU, which is within the range of measurements observed

at G312. Therefore, low pH conditions at G312 are attributable to natural variability within the aquifer.

4.5 LOE #5: pH values within the proposed monitoring well network are strongly correlated with saturation indices of carbonate minerals in soil near AP1

As mentioned in Section 4.3, composite soil samples from various locations surrounding AP1 were collected and submitted for mineral identification analysis via XRD (**Table 2**). Soil surrounding AP1 contains variable abundances of carbonate minerals such as calcite, dolomite, and ankerite, with the total abundance of carbonates at each location ranging an order of magnitude from 2.7-27.5 wt.%. Carbonate minerals in nature function as pH buffers, capable of neutralizing acidity through reaction with carbonate (CO_3) (Drever, 1988). pH levels at individual wells may be significantly influenced by the presence and abundance of carbonate minerals comprising localized sections of the aquifer unit. Although soil samples were not collected for all wells of interest, carbonate saturation indices (SIs) provide a method to assess the role of carbonate minerals in soil buffering capacity in the absence of XRD results.

United States Geologic Survey (USGS) software package PHREEQC was used to calculate SIs of carbonate minerals at G312 and background wells G281 and G306 based on groundwater compositions. A mineral's SI is an expression of its thermodynamic equilibrium state relative to a liquid (groundwater). If the calculated SI for a mineral is negative, then that mineral is undersaturated relative to groundwater and is thermodynamically favored to dissolve. If the calculated SI for a mineral is positive, then that mineral is supersaturated relative to groundwater and is thermodynamically favored to precipitate. If a mineral's SI is approximately $0 (+10.2)$, then the mineral is in thermodynamic equilibrium with groundwater. SIs for calcite $(CaCO₃)$ and dolomite (Ca,Mg(CO3)2) were plotted against pH for individual samples (**Figure 5**). **Figure 5** demonstrates a strong positive correlation between pH and carbonate SI. pH values tend to be lower in groundwater that is undersaturated with respect to carbonate minerals. This relationship is expected – monitoring wells which favor carbonate dissolution are likely to contain less carbonate in the solid phase. Absence of carbonate in localized portions of the aquifer results in the inability of these locations to buffer low pH groundwater. According to **Figure 5**, background wells G281 and G306 are near equilibrium or supersaturated with respect to carbonate minerals and are likely to have these minerals present and stable. These wells would then have greater capability to buffer acidic water and retain near-neutral pH values. G312 was not sampled for mineralogy, although **Figure 5** demonstrates that groundwater from this well is undersaturated with respect to carbonate minerals, suggesting that large abundances of carbonate are not likely to be present in aquifer solids at this location. tioned in Section 4.3, composite soil samples from various locations surrounding AH
d and submitted for mineral identification analysis via XRD (Table 2). Soil surround
tatains variable abundance of carbonates minreals su

XRD analyses indicate carbonate mineral abundances around AP1 vary up to an order of magnitude (**Table 2**). Evaluation of carbonate SIs reveals that a strong correlation exists between

carbonate SIs and pH. G312 is undersaturated with respect to calcite and dolomite; therefore, these minerals are likely not present as pH buffers, resulting in lower groundwater pH values where acid neutralizing minerals are not available.

RAFT

SECTION 5

CONCLUSIONS

Based on these five LOEs, it has been demonstrated that AP1 is not the source of the potential cobalt and pH exceedances identified.

- 1. AP1 porewater samples do not contain detectable concentrations of cobalt, whereas cobalt concentrations in background well G306 occasionally exceed the relevant GWPS.
- 2. Cobalt concentrations in ash samples collected from AP1 are comparable to or lower than cobalt concentrations in soil samples from downgradient and background monitoring wells.
- 3. Monitoring well G314 has experienced significant changes in oxidation-reduction (redox) conditions since well installation occurred, which may cause destabilization of ironbearing minerals capable of hosting cobalt ions in their crystal structure. AP1 porewater samples do not contain detectable concentrations of cobalt, whereasement
rations in background well G306 occasionally exceed the relevant GWPS.
Cobalt concentrations in ash samples collected from Δ P1 are c
- 4. AP1 porewater is slightly basic and would not result in low pH measurements at monitoring well G312.
- 5. pH values within the proposed monitoring well network are strongly correlated with saturation indices of carbonate minerals which are detected at variable abundances across soil near AP1.

SECTION 6

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- Ramboll. 2021b. Hydrogeologic Site Characterization Report, Ash Pond No. 1, Coffeen Power Plant, Coffeen, Illinois. October.

TABLES TABLES

Table 1: Cobalt Concentrations in Soil and Ash Coffeen Power Plant - Ash Pond No. 1

Notes:

Soil samples were composite samples collected over the indicated depth range The composite samples collected over the indicated depth range

Table 2: Summary of X-ray Diffraction Results *Geosyntec Consultants, Inc.* **Coffeen Power Plant - Ash Pond No. 1**

Notes:

Results presented in units of weight %

- : Mineral was not detected in sample

Weight % quantities have been normalized to a sum of 100% to remove reporting of amorphous material Carbonate total consists of calcite, dolomite, and ankerite

FIGURES FIGURES

water samples XPW01, XPW02, and AP1d. The calculated Groundwater Protection Standard (GWPS) is indicated by the dashed line. Samples which did not contain cobalt concentrations above the method detection limit of 0.002 mg/L are displayed on the figure as having a detected concentration of 0.002 mg/L.

APPENDIX A

Figure 2-1: Proposed 845 Groundwater Monitoring Well Network. From Groundwater Monitoring Plan, Ash Pond No. 1, Coffeen Power Plant APPENDIX A

1: Proposed 845 Groundwater Monitor

rk. From Groundwater Monitoring Pla

Pond No. 1, Coffeen Power Plant

PROPOSED 845 GROUNDWATER MONITORING WELL NETWORK

FIGURE 2-1

RAMBOLL AMERICAS ENGINEERING SOLUTIONS, INC.

 \Box COMPLIANCE WELL

STAFF GAGE

DE BACKGROUND WELL SITE FEATURE PART 845 REGULATED UNIT (SUBJECT UNIT) LIMITS OF FINAL COVER **CONTROPERTY BOUNDARY**

GROUNDWATER MONITORING PLAN ASH POND NO.1 COFFEEN POWER PLANT COFFEEN, ILLINOIS

APPENDIX B

Figure 1-3: Uppermost Aquifer Groundwater Elevation Contours, April 20, 2021. From Groundwater Monitoring Plan, Ash Pond No. 1, Coffeen Power Plant APPENDIX B
3: Uppermost Aquifer Groundwater E
tours, April 20, 2021. From Groundwa
ing Plan, Ash Pond No. 1, Coffeen Pow

PROJECT: 169000XXXX | DATED: 10/12/2021 | DESIGNER: HOTCA

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0 275 550 Feet \perp

(SUBJECT UNIT) SITE FEATURE

STAFF GAGE

 \bigodot

PROPERTY BOUNDARY

 \blacksquare MONITORING WELL

PART 845 REGULATED UNIT LIMITS OF FINAL COVER GROUNDWATER ELEVATION CONTOUR (2-FT CONTOUR INTERVAL, NAVD88) INFERRED GROUNDWATER ELEVATION CONTOUR **GROUNDWATER FLOW** DIRECTION

FIGURE 1-3

RAMBOLL AMERICAS ENGINEERING SOLUTIONS, INC.

UPPERMOST AQUIFER GROUNDWATER ELEVATION CONTOURS APRIL 20, 2021

NOTE:

ELEVATIONS IN PARENTHESES WERE NOT USED FOR CONTOURING.

APPENDIX B MODFLOW, MT3DMS, and HELP MODEL FILES (ELECTRONIC ONLY) FLOW, MT3DMS, and HELP MODEL FILES
CTRONIC ONLY)

APPENDIX C EVALUATION OF PARTITION COEFFICIENT RESULTS (GEOSYNTEC CONSULTANTS, INC., 2022) ENDIX C
LUATION OF PARTITION COEFFICIENT RESULTS
SYNTEC CONSULTANTS, INC., 2022)

Memorandum

INTRODUCTION

The Illinois Power Generation Company (IPGC) currently operates the Coffeen Power Plant (CPP) and its associated ash ponds located in Coffeen, Illinois. Ash Pond Number (No.) 1 (AP1) (Vistra identification (ID) No. 101; Illinois Environmental Protection Agency [IEPA] ID No. W1350150004-01; National Inventory of Dams [NID] No. IL50722) is a 23-acre, unlined SI used to manage CCR (bottom ash) and non-CCR waste streams at the CPP in accordance with the plant's Water Pollution Control Permit 1978-EA-389 issued by the Agency on May 26, 1978. Geosyntec Consultants (Geosyntec) is assisting IPGC with Part 845 compliance at the Site.

IPGC is currently preparing a Construction Permit application for AP1 as required under Section 845.220. As part of the Construction Permit application, groundwater modeling is being completed for known potential exceedances of groundwater protection standards (GWPS) identified in the Operating Permit (Burns & McDonnell, 2021). In the Operating Permit (October 2021), Burns & McDonnell identified potential GWPS exceedances for several compounds potentially associated with AP1, including boron, cobalt, pH (field), sulfate, and total dissolved solids (TDS). An evaluation of potential exceedances of applicable GWPS found that both cobalt and pH potential exceedances are not related to AP1 (Geosyntec, 2022). Batch adsorption testing was conducted for boron and sulfate to generate site-specific partition coefficients. This technical memorandum summarizes the results of the batch adsorption testing and calculation of partition coefficients. sto: Brian Hennings - Ramboll

Allison Kreinberg, Ryan Fimmen - Geosyntee Consultants, Inc.

11. Draft Evaluation of Partition Coefficient Results - Coffieen Ash Pond No

CCR Unit 101, Colfieen Power Plant, Colfieen, Illin

IPGC – AP1 Batch Attenuation Testing Summary May 11, 2022 Page 2

BATCH ATTENUATION TESTING

In 2021, Geosyntec conducted a field investigation at AP1 which included completion of four (4) soil/rock borings ranging in depth from 13 to 18 feet below ground surface. As part of that investigation, soil and groundwater samples were submitted to SiREM Laboratories (Guelph, ON) for batch solid/liquid partitioning testing. A summary of the soil samples used for the batch testing is provided in **Table 1**.

Two groundwater samples (G311 and G313) and three soil samples (SB-306, SB-311, and SB-313) were used for batch attenuation testing at five (5) soil:solution ratios (**Table 1**), each ran in duplicate. For each treatment, 0.1 L of groundwater was brought into contact with varying amounts of soil (0.004 to 0.2 kg, depending on the ratio) and equilibrated over a seven-day period. Each microcosm was amended (i.e., spiked) with sodium sulfate (Na2SO4), and the microcosms with G313 groundwater were also amended with boric acid (H₃BO₃), to achieve a target concentration of sulfate and boron, respectively (**Table 2**). The G311 microcosm was not amended with boric acid because potential boron exceedances were not identified in the vicinity of G311. G313 groundwater was combined with aquifer solids both adjacent to downgradient location G311 and background location G306 to understand how partitioning behavior may be affected by position relative to AP1. is sometimal plantationing usamig. A saminality of the soil samples sises to the band of solid Table 1.

ded in Table 1.

condwater samples ((3311 and (3313) and three soil samples (813-306, 813-311, a

condwater samples

An initial sample of the stock solution for each experimental design was collected on Day 0, and a control sample (i.e., only amended G311 or G313 groundwater with no aquifer solids) was collected on Day 7 after tumbling in polypropylene bottleware to evaluate any loss to interactions with the bottleware or ambient conditions. Duplicates were constructed for each microcosm, including the control samples. After seven days of contact time, an aliquot of the free liquid was collected and filtered through a 0.45 micron (μm) filter prior to analysis for dissolved concentrations of sulfate and/or boron. The oxidation/reduction potential (redox) and pH were measured for each batch test at the beginning and end of the contact period and in the control samples.

Data obtained from the tests (**Tables 3** and **4**) were used to construct isotherms for boron and sulfate; 5-point isotherms were constructed by averaging duplicate results for each soil:solution ratio. Mathematical fitting was used to calculate the attenuation distribution coefficients (K_d) , assuming linear adsorption. The linear adsorption equation was used:

$$
q_e = K_d \times C_e \tag{Eq. 1}
$$

where q_e is the mass of constituent adsorbed to the solid phase at equilibrium, C_e is the remaining aqueous constituent concentration at equilibrium, and K_d is the linear sorption coefficient (reported in liters per kilogram [L/kg]). Some of the data showed a deviation from a linear trend, and so were also fitted using non-linear isotherms. The non-linear Langmuir isotherm was used:

IPGC – AP1 Batch Attenuation Testing Summary May 11, 2022 Page 3

$$
q_e = \frac{q_m K_L C_e}{1 + K_L C_e} \tag{Eq. 2}
$$

where *qm* is the inverse of the slope and *KL* is the Langmuir distribution coefficient. The adsorption data were linearized according to:

$$
\frac{C_e}{q_e} = \frac{1}{(K_L \times q_m)} + \frac{C_e}{q_m}
$$
 Eq. 3

A common non-linear Freundlich equation was also used:

$$
q_e = K_F (C_e)^{1/n}
$$
 Eq. 4

where q_e is the mass of constituent adsorbed to the solid phase at equilibrium, C_e is the remaining aqueous constituent concentration at equilibrium, *KF* is the Freundlich distribution coefficient, and *1/n* is a non-linearity constant. The adsorption data were plotted as log-transformed values to perform the non-linear isotherm fitting using the linearized Freundlich equation:

$$
log(q_e) = log(K_F) + (1/n)log(C_e)
$$
 Eq. 5

The calculated linear, Langmuir, and Freundlich distribution coefficients (K_d, K_L, and K_F, respectively) and 1/*n* values are shown in **Tables 5** and **6**.

SUMMARY OF RESULTS

The partition coefficient values for G311 and G313 (denoted below as G313/SB-306 when combined with SB-306 geologic material and G313/SB-313 when combined with the SB-313 geologic material) are presented in **Tables 5** and **6**, respectively. Figures which show the linear, Langmuir, and Freundlich isotherms for boron and sulfate are provided in **Appendix A**.

A boron partition coefficient was not calculated for G311, since the microcosm was not amended with boric acid because potential boron exceedances were not identified in the vicinity. The Freundlich isotherm fit the data best for G313/SB-306 and G313/SB-313, yielding KF values of 0.65 L/kg and 2.03 L/kg, respectively. Though slightly higher at G313/SB-313, these values are comparable to boron partition coefficients reported in the literature, which range from 0.19 to 1.3 L/kg depending on pH conditions and the amount of sorbent present (EPRI, 2005; Strenge & Peterson, 1989). $\frac{R_2}{R_2} = \frac{1}{(R_2 \times q_{\rm m})} + \frac{C_2}{q_{\rm m}}$
 $\frac{1}{R_2} = \frac{1}{(R_2 \times q_{\rm m})} + \frac{C_2}{q_{\rm m}}$

(i.e. $\frac{1}{(R_2 \times q_{\rm m})} + \frac{C_2}{q_{\rm m}}$ is the meass of constituent adsorbed to the solid phase at equilibrium, C_c is the re

The G311 partition coefficient for sulfate ranged from -624 L/kg for the Langmuir isotherm to 10.11 L/kg for the linear isotherm, but the best-fitting Freundlich isotherm yielded a low K_F value of 9.2 \times 10⁻¹² L/kg. None of the isotherms showed a high goodness-of-fit (i.e., R²) for either G313/SB-306 or G313/SB-313, with the highest correlation being 0.51, and were associated with erroneously high (1700 L/kg) and low (-690 L/kg) partition coefficients. An accurate sulfate

IPGC – AP1 Batch Attenuation Testing Summary May 11, 2022 Page 4

partition coefficient could therefore not be calculated from any of the data. These results are consistent with the findings of Strenge and Peterson (1989), who found that partition coefficients for sulfate are 0.0 L/kg, regardless of pH conditions and the amount of sorbent present.

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TABLES TABLES

Table 1 - Batch Attenuation Testing Data Summary Coffeen AP1 *Geosyntec Consultants*

Notes:

ft bgs = feet below ground surface
Table 2 - Microcosm Amendment and Target Concentrations Coffeen AP1

Notes:

ft bgs - feet below ground surface

mg/L - milligrams per liter

 $Na₂SO₄$ - sodium sulfate

 H_3BO_3 - boric acid

Table 3 - Batch Attenuation Testing Results, G311 Coffeen AP1

Geosyntec Consultants

Notes:

mg/L - milligrams per liter

mV - millivolts

SU - Standard Units

ORP - oxidation/reduction potential

Table 4 - Batch Attenuation Testing Results, G313 Coffeen AP1

Geosyntec Consultants

Notes:

mg/L - milligrams per liter

mV - millivolts

SU - Standard Units

ORP - oxidation/reduction potential

Table 5 - Partition Coefficient Results, G311 Coffeen AP1

Notes:

 $\rm K_D$ - linear partition coefficient

KL - Langmuir partition coefficient

 \mathbf{K}_{F} - Freundlich partition coefficient

 \mathbf{q}_m - inverse of the slope of the linearized Langmuir isotherm

Table 6 - Partition Coefficient Results, G313 Coffeen AP1

Notes:

The Freundlich isotherm was not calculated for G313/SB-313

because the data were not conducive to log transformation

 K_D - linear partition coefficient

KL - Langmuir partition coefficient

 \mathbf{K}_{F} - Freundlich partition coefficient

qm - inverse of the slope of the linearized Langmuir isotherm

APPENDIX A BATCH TESTING ISOTHERM PLOTS APPENDIX A
BATCH TESTING ISOTHERM PLOTS

